

INTRODUCTION
The Montana Bureau of Mines and Geology (MBMG), in conjunction with the STATEMAP advisory committee, selected the Argenta 7.5' quadrangle in southwest Montana for detailed mapping as part of the MBMG's ongoing effort to complete the Dillon 30' x 60' (scale 1:100,000) geologic map. A key goal of the mapping was to reevaluate the Mesoproterozoic rocks in the northwest part of the quadrangle using new stratigraphic interpretations for the southern Belt Basin developed by the MBMG and the Idaho Geological Survey in the west-adjacent Salmon 30' x 60' quadrangle (Lonn and others, 2019).

The Argenta quadrangle is located northwest of Dillon, Montana and encompasses the mountains and foothills at the southeastern end of the Pioneer Mountains (Fig. 1). The quadrangle lies within the Argenta Mining District where gold, silver, copper, lead, and zinc were produced from lode and placer deposits (Gusch, 1972). The best access is via gravel roads along Rattlesnake Creek. The majority of the quadrangle is public land administered by the U.S. Forest Service and the Bureau of Land Management.

PREVIOUS MAPPING AND METHODS
The Argenta quadrangle is included in small-scale mapping by Ruppel and others (1993, scale 1:250,000) and larger-scale mapping by Myers (1952, scale 1:31,680) and Hobbs (1967, scale 1:24,000; Fig. 2). Unpublished mapping by Alan English (MBMG) and the University of Montana Western geology field camp students, under the direction of Dr. Robert Thomas, was completed for parts of this map. New field mapping was completed in 2015, using a U.S. Geological Survey 1:24,000-scale topographic base map, 2009 orthorectified from the National Agricultural Imagery Program (NAIP), and a handheld Trimble Juno GPS for locating sample and field observation point data. Rock samples collected for U-Pb geochronology were processed at the MBMG mineral separation laboratory. Zircon was isolated from selected samples by standard density and magnetic separation techniques. Zircon separates were analyzed by laser ablation inductively coupled plasma-mass spectrometry (LA-ICP-MS) by Jesse Mosoff, MBMG, at the University of California, Santa Barbara.

GEOLOGIC SUMMARY
The oldest rocks in the quadrangle are exposed in the core of the Humboldt Mountain Anticline (Fig. 1). These sedimentary units were previously interpreted as part of the Mesoproterozoic Belt Supergroup (Myers, 1952; Ruppel and others, 1993), however, we interpret them as Cambrian and/or Neoproterozoic and/or Mesoproterozoic (map unit C2YAR). We assign the uppermost part of these rocks (map unit Cq), consisting of light gray quartz arenite with discontinuous interbedded argillite and siltstone, as Cambrian based on the discovery of abundant trace fossils (Fig. 3) in argillite at the top of Humboldt Mountain. This uppermost unit may correlate with the basal Cambrian Flathead Formation, which is generally considered Middle Cambrian. However, the reported occurrence of Trepostichus pedum (Tripp, 2019), the trace fossil whose first appearance defines the Precambrian-Cambrian boundary (Brasier and others, 1994), suggests this unit could be Early Cambrian.

Underlying the gray quartzites and dark argillites are poorly exposed, predominantly red and green micaceous argillites and interbedded reddish quartz sandstone and siltstone (map unit C2YAR). We tentatively interpret this unit as Cambrian, Neoproterozoic, and/or Mesoproterozoic based on its stratigraphic position below the Precambrian-Cambrian boundary (Brasier and others, 1994), which also suggests parts of it could be Neoproterozoic or younger.

The majority of the quadrangle is underlain by Paleozoic and Mesozoic carbonate and clastic rocks exposed in a series of northeast-trending folds. Cretaceous igneous rocks, including small intrusions and minor volcanic flows associated with the Pioneer Batholith, intrude and overlie the folded sedimentary strata. Cenozoic volcanic and sedimentary deposits are widespread in the southern part of the map. The volcanic rocks include rhyolite tuffs and isolated andesitic flows that are part of the Eocene Dilton Volcanic Group (Fritz and others, 2007). These volcanic rocks erupted during onset of regional Eocene extension in southwest Montana (e.g., Foster and others, 2010), and are overlain by middle Eocene to early Miocene sedimentary deposits of the Renova Formation.

In the central part of the map, north of Rattlesnake Creek, cobble conglomerate within the Renova Formation (map unit Tre) represents fluvial basin-margin facies associated with Paleogene drainage basins. The fine-grained, sub-arkic sandstones (map unit Tr) reflect basin interior facies (Harber and others, 2012). Younger cobble and boulder gravel deposits dominated by rounded quartzite clasts are interpreted as alluvial-fan and debris flow deposits correlated with the Miocene to Pliocene Sixmile Creek Formation (Thomas and Sears, 2020).

The youngest deposits in the quadrangle are Quaternary surficial deposits. Argenta Flats is underlain by a large Pleistocene glacial outwash fan that is an important aquifer in the Beaverhead Valley (Thomas and others, 2007). Extensive quartzite talus and colluvium mantle the slopes of Humboldt and Dutchman mountains (Fig. 1). Alluvium and colluvium occur throughout the map area.

STRUCTURE
The oldest structures in the Argenta quadrangle are northeast-trending folds and thrust faults associated with Late Cretaceous crustal shortening at the leading edge of the Cordilleran fold-thrust belt of southwest Montana (Ruppel and Lopez, 1984). The folds, including the Humboldt and Dutchman Mountain Anticlines and the Cave Gulch Syncline, range from upright to overturned and generally plunge to the north. The thrust faults also trend northeast, dip to the west, and locally cut the anticlines. The thrust faults in the northwest corner of the map was interpreted by Ruppel and others (1993) as a klippe related to the nearby Grasshopper Thrust (Fig. 1, Kelley Thrust of Myers, 1952). Our interpretation, based partly on the recognition of previously unmapped Cambrian rock around Humboldt Mountain, shows it as an imbricate thrust in the footwall of the Grasshopper Thrust, similar to the other mapped thrusts in the quadrangle.

The youngest structures are steep, generally northwest and northeast-trending Cenozoic Basin and Range extensional faults that offset the older structures. Rare, poorly exposed fault surfaces include both dip-slip and strike-slip motion. No Quaternary faults have been mapped in the quadrangle but the epicenter of the 2005 Mw 5.6 Dillon earthquake (Stockey, 2007) was located approximately 11 km (7 mi) westward of the quadrangle (Fig. 1).

DESCRIPTION OF MAP UNITS

Qal Alluvium (Holocene)—Unconsolidated gravel, sand, silt, and clay deposited by Rattlesnake Creek and small tributaries. Clasts generally subrounded to well-sorted and consist predominantly of quartzite. Thickness as much as 10 m (33 ft).

Qac Alluvium and colluvium (Holocene and Pleistocene)—Unconsolidated sand, silt, and clay, and subordinate gravel, deposited on gentle slopes by sheetwash and gravity processes. Variable thickness, generally less than 10 m (33 ft).

Qc Colluvium (Holocene and Pleistocene)—Unconsolidated, locally derived slope deposit that contains angular, poorly size-sorted clasts, generally pebble size and larger. Thickness less than 10 m (33 ft).

Qta Talus deposit (Holocene and Pleistocene)—Unconsolidated, locally derived aprone-like deposit with angular clasts on and below steep slopes. Includes some rock-slope deposits. Variable thickness, typically less than 10 m (33 ft).

Qal Alluvial-fan deposit (Holocene and Pleistocene)—Unconsolidated, poorly sorted cobbles, gravel, sand, and silt forming fan-shaped deposits. Thickness probably less than 6 m (20 ft).

Qdf Debris-flow deposit (Holocene and Pleistocene)—Angular, subangular, and subrounded clasts of chaotic and unstratified boulders and cobbles and subordinate fine-grained sediment. Some fine sediment probably has been removed by erosion leaving coarser clasts as lag deposits. Up to approximately 15 m (50 ft) thick.

Qgr Gravel deposit (Holocene and Pleistocene)—Poorly sorted gravel, sand, and silt with angular to well-rounded cobbles that consist dominantly of quartzite and limestone. Occurs as isolated deposits in and along smaller drainages and includes alluvium, small debris flows, alluvial fan deposits, and colluvium. Thickness up to 5 m (15 ft).

Qcg Pediment gravel (Holocene and Pleistocene)—Thin accumulations of cobbles, pebbles, and scattered small boulders with a sand and silt matrix capping low-relief, eroded bedrock surfaces. Light brown weathared quartzite cobbles that appear to be the Quadrant Formation are most abundant. Occurs as a cobble lag deposit in some places. Thickness as much as 5 m (15 ft).

Qgo Glacial outwash deposit (Pleistocene)—Weakly stratified deposit of sand and gravel with abundant, locally derived, rounded to subrounded cobbles and boulders. Quartzite and feldspathic quartzite clasts are mostly flat, angular, granitic clasts, argillite, hornfels, limestone, sandstone, and volcanic clasts are less abundant. Forms an extensive outwash fan deposit that underlies Argenta Flats. Also occurs as an isolated deposit overlying bedrock north of Rattlesnake Creek near the western edge of the map. Domestic well logs indicate the gravels are as much as 13 m (42 ft) thick and are underlain by up to 90 m (295 ft) of clay-rich gravel and cobble layers that may be an outwash.

Tsc Gravel and conglomerate deposits (Pleistocene)—Unconsolidated, poorly sorted, massive to crudely bedded, gravel and conglomerate in a matrix of sand, silt, and clay. Clasts are subangular to subrounded cobbles and boulders (some up to 2 m in diameter) consisting dominantly of quartzite and subordinate sandstone, argillite, limestone, chert, volcanic, and plutonic rock. Unconformably overlies Eocene rhyolite tuff (map unit Tr) in southeast corner of the map. May be equivalent to the Miocene Sweetwater Creek member of the Sixmile Creek Formation, which has been interpreted as debris flows on alluvial fans (Thomas and Sears, 2020). Thickness as much as 90 m (300 ft).

Tec Conglomerate and sandstone (Tertiary; Oligocene)—Angular to subrounded pebbles and cobbles in a poorly sorted, silt to granule, calcareous matrix. This is a thick bed with irregular contacts are generally massive with rare normal to reverse grading. Clasts, up to 35 cm (14 in), are moderately indurated and consist of light to dark gray, blocky, angular to subangular limestone (55 percent), sandstone and siltstone (21 percent), quartz sandstone (12 percent), feldspathic quartzite (8 percent), tuff (4 percent), and chert (4 percent) (Barber, 2013). This unit has been interpreted by Barber (2013) as proximal debris-flow-dominated, alluvial fan deposits. The youngest grain from a sample collected for detrital zircon analysis (FP13-4, Schwartz and others, 2019) yielded a U-Pb age of 25.2 ± 0.1 Ma, suggesting Oligocene deposition. Base not exposed but as much as about 90–120 m (300–400 ft) thick.

Tre Renova Formation (Tertiary; Oligocene–Eocene)—Very pale orange, grayish orange, white, and pale yellowish brown, calcareous, medium- to coarse-grained, light- to fine-grained sandstone. Sandstone intervals are poorly sorted, light micaceous, and angular to subangular granitic pebbles and cobbles. Carbonate clasts are predominantly quartzite, limestone, and volcanic rock. Locally contains petrified wood, paper shales, and clay-rich intervals that we refer to as "popcorn" texture. Estimated thickness around 790 m (2,600 ft) thick.

Trt Rhyolite tuff (Tertiary; Eocene)—Massive, white, weathering, vitric tuff. Contains broken phenocrysts (up to 30 percent) of alkali feldspar, quartz, biotite, and occasional amphibole in a vitric groundmass composed of cuspidate glass shards that are mostly fresh, but are locally desiccated. Tuff intervals usually contain small fragments of pumice, volcanic debris, fine-grained sedimentary rock, and silicified wood fragments. Thin intervals of plane, low-angle cross bedding is observed locally. This unit yielded a U-Pb age of 41.2 ± 0.2 Ma (table 1, sample DN-AR1). Base not exposed but at least 60 m (200 ft) thick.

Taf Andesitic flows (Tertiary; Eocene)—Dark brown to black andesitic flows with an aphanitic groundmass and 2–3 percent phenocrysts of olivine, pyroxene, rare plagioclase, and quartz. Forms isolated, knobby outcrops that weather pale reddish brown to black and often have a vesicular texture.

Kgd Granodiorite (Late Cretaceous)—Medium- and fine-grained, light to medium gray, granodiorite and quartz monzonite. Hornblende and biotite are the dominant mafic minerals. Occurs as several isolated bodies that may be satellite intrusions to the Uphill Creek Pluton (Snee, 1978, 1982) of the Pioneer Batholith. This unit yielded a U-Pb age of 71.0 ± 0.5 Ma (table 1, sample DN-AR2), similar to an "Ar" biotite age of 71.5 ± 0.6 Ma reported by Snee (1982) for a sample collected near the town of Argenta.

Kov Cold Spring Creek volcanics (Late Cretaceous)—Variegated, intercalated sequence of porphyritic lava flows and ash-flow tuff. Named for exposures along Cold Spring Creek (southeast of quadrangle) described by Ivy (1988). Porphyritic lavas are generally dark colored with phenocrysts of plagioclase, orthopyroxene, clinopyroxene, hornblende, biotite, and quartz in an aphanitic groundmass. Poorly exposed tuffs contain varying amounts of pumice fragments and accidental clasts of porphyry lava. Ivy (1988) reported "Ar" hornblende ages spanning ~80–76 Ma for this unit. U-Pb zircon ages from nearby mapping (Mosoff, 2021; Mosoff, in prep.) ranged from 71.7 ± 0.6 and 73.1 ± 0.8 Ma. Thickness as much as 490 m (1,600 ft) south of the map area.

Kf Frontier Formation (Late Cretaceous)—Dominantly gray, brownish gray, and greenish gray siltstone and mudstone, and subordinate medium- to coarse-grained, and locally very coarse-grained sandstone, conglomerate, limestone, and minor porcellanite. The beds of mudstone, siltstone, limestone, and sandstone form fine-upgrading depositional cycles tens of meters thick. Sandstones and conglomerates are rich in quartz and chert. Conglomerate clasts are rounded pebbles and small cobbles. Volcaniclastic sandstone and bentonitic mudstone occur in upper part. Lower 100–200 m (330–660 ft) is distinctive brown to brownish gray siltstone and mudstone. Top not exposed but thickness is about 900 m (3,000 ft) north of the quadrangle (Tysdal and others, 1994; Dymann and Tysdal, 1998).

Kov Blackfoot Formation, Vaughn Member (Early to Late Cretaceous)—Olive green, yellowish green, bright green, and gray green, hard, dense, and calcareous siltstone and porcellanite (silicified) mudstone. Subordinate gray, greenish gray and olive gray, fine- to medium-grained, and locally coarse-grained siltstone, with high percentage of volcaniclastic debris, and matrix-supported conglomerate and conglomeratic sandstone with clasts mostly of chert and quartzite. An association of distinctive lithologies is present in the uppermost part, which consists of massive mudstone and siltstone, gray freshwater limestone or locally very calcareous mudstone and siltstone, dark gray shale, and bright green porcellanite. Upper contact is mapped at top of a porcellanite bed, interbedded with micritic limestone, that directly overlies the highest maroon mudstone (Tysdal and others, 1994; Dymann and Tysdal, 1998). A porcellanite bed approximately 25 m (80 ft) below the top of the Vaughn in the Fryling Fan Basin yielded a U-Pb age of 94.8 ± 0.5 Ma (Zartman and others, 1995). Thickness approximately 584 m (1,916 ft).

Kov Blackfoot Formation, Flood Member (Early Cretaceous)—Upper part: pale brown to brownish gray, fine- to medium-grained and locally coarse-grained to conglomeratic, quartz- and chert-rich siltstone, and conglomerate. Trough crossbedding common in sandstone. Middle part dominantly gray mudstone, shale, and minor interbeds of siltstone and quartz-rich sandstone. Lower part: medium gray and locally green and red calcareous siltstone and mudstone, gray shale, and gray calcareous, fine- to medium-grained sandstone that is rich in quartz and chert grains (Tysdal and others, 1994; Dymann and Nichols, 1988). Thickness 183–213 m (600–700 ft).

Kov Kootenai Formation (Early Cretaceous)—Mapped as one unit, but consists of four distinct units (after Myers, 1952). Combined thickness of all units is approximately 274–305 m (900–1000 ft) thick. The uppermost unit is the fine-grained sandstone *Gastropod limestone member*. Light gray, thick-bedded, gastropod coquina or gastropod-rich limestone that may also contain chrysolites and ostracodes. Forms conspicuous ridges.

Kov *Red mudstone member*: Variegated shale and mudstone, dominated by red, orange, and purple, and subordinate light and medium gray colors, interbedded with minor reddish quartzite and chert-rich limestones. Poorly exposed, recessive unit.

Kov *Fine-grained limestone member*: Pale yellowish-gray to pale brown, dense limestone and shale limestone with interbedded shale.

Kov *Basal sandstone and mudstone*: Upper part recessive, mostly reddish and greenish mudstone; lower part is ridge-forming, coarse- to medium-grained, cross-bedded to massive, brown to yellowish gray chert-rich limestones with local lenses of chert-rich conglomerate and limestone pebble conglomerate; interbedded with reddish and greenish mudstone.

Jm Morrison Formation (Jurassic)—Green, red, and gray variegated mudstone, shale, and siltstone with thin, interbedded yellowish brown to grayish shales, very fine-grained sandstone and siltstone beds, and thin, gray limestone beds. Thickness about 45–55 m (150–175 ft).

Kd Dinwoody Formation (Lower Triassic)—Interbedded shale, limestone, and calcareous sandstone characterized by distinctive pale to light gray brown weathered surfaces and the presence of (possibly pelecypod) Lingulid. Upper part has massive calcareous, rippled sandstone beds as much as 1 m (3 ft) thick with shaly interbeds. Massive, gray, pinkish gray weathering limestone as much as 1 m (3 ft) thick is a major component in the upper part. Lower part is predominantly olive drab, chippy-weathering, fair to fine shale with interbedded dark brown weathering, silty limestone beds 10 cm (4 in) or thicker. Thickness about 244 m (800 ft).

Psp Sheldahl and Phosphoria Formations, undivided (Permian)—Upper part is brownish gray to dark gray, laminated or thin- to thick-bedded chert, gray brown quartz sandstone, and cherty sandstone. The mostly fine-grained, well-sorted quartz sandstone occurs on cylindrical burrows and is interpreted as the Sheldahl Formation described by McKelvey and others (1959). Lower part of the unit is dark gray to black, carbonaceous and phosphatic mudstone with scarce porphyritic beds, grayish and gray-brown cherty quartz sandstone, cherty sandy dolomite, fine-grained dolomitic sandstone, and yellowish tan sandy siltstone with subordinate beds of vitric tuff sandstone. Poorly exposed, typically covered by colluvium, and overlies underlying Quadrant Formation. Thickness approximately 183–215 m (600–700 ft).

Pq Quadrant Formation (Pennsylvanian)—Light gray to light yellowish brown, fine-grained, vitreous quartz sandstone. Beds are mostly thick to massive, occasionally with faint cross-lamination. Locally, the basal quartz sandstone beds are interbedded with intervals of limestone or dolomite and locally have gray limestone pebbly clasts. Forms resistant ridges typically covered with conifers. Thickness approximately 244 m (800 ft).

Pfmr Snowcrest Range Group (Pennsylvanian–Mississippian)—Originally mapped as Anson Formation (Myers, 1952; Hobbs, 1967) but shown here as the Snowcrest Range Group, which includes the Conover Ranch, Lombard, and Kibby Formations (Wardlaw and Peacor, 1985). Total thickness of Snowcrest Range Group is about 238 m (780 ft) but localized outcrop-scale folds, especially in the Lombard, make thickness estimates problematic.

Pfmr *Conover Ranch Formation*: Pale reddish and pale yellowish, thin-bedded, calcareous mudstone with minor interbeds of limestone, calcareous sandstone, and siltstone. Some beds contain brachiopods, bryozoan corals, crinoid columns, and belemnite fragments.

Pfmr *Lombard Formation*: Upper unit is fossiliferous, thin- to thick-bedded, lime-mudstone and limestone with rare chert lenses. Brachiopods, rugose corals, corals, and small conical lenses are common. Lower unit is poorly fossiliferous, thin to thick, indistinctly bedded, lime-mudstone and gray limestone.

Pfmr *Kibby Formation*: Gray quartzite and yellowish, calcareous sandstone beds interbedded with siltstone and shale.

Mmc Mission Canyon Formation (Mississippian)—Light gray, thick-bedded, fossiliferous, often chert-forming limestone with irregular chert bands. Myers (1952) mapped the Mission Canyon in two parts that were combined for this map. The upper part is light to dark gray, commonly well-bedded, lime- to medium-grained, with rare coarse-grained beds of crinoidal limestone, and minor light to dark gray, thin-bedded dolomite, locally sandy and cherty limestone. The lower part is generally a coarser grained limestone with abundant crinoids and some intervals of thin-bedded limestone and dolomitic limestone. Becomes coarsely crystalline marble near intrusions. Thickness variable as a result of deformation but estimated to be about 360 m (1,175 ft) thick.

Ml Lodgepole Formation (Mississippian)—Gray to dark gray, thin- to medium-bedded, fossiliferous limestone. Thinner beds and shaly partings are common near base, beds are thicker with rare shaly partings near the top. Common fossils are corals, crinoids, and rare brachiopods. The Lodgepole in this area has less chert than in other areas of Montana to the north and east. Becomes coarsely crystalline marble near intrusions. Thickness variable as a result of deformation but generally about 290 m (950 ft) thick.

MDI Three Forks and Jefferson Formations, undivided (Mississippian and Devonian)

MDI *Three Forks Formation* (Mississippian and Devonian)—Brown, argillaceous, fossiliferous limestone interlayered with black to dark gray, carbonaceous shale, grayish green shale, and light tan, silty sandstone. Recessive and mapped on the basis of metamorphosed to hornfels near intrusions. Thickness approximately 75 m (250 ft).

D Jefferson Formation (Devonian)—Upper part is yellowish-weathered solution breccia of fine-grained calcite marble. Lower part is medium- to thick-bedded, dark gray to brown, coarsely crystalline, weakly mottled dolomite with strong porphyroblasts and fresh surfaces. Outcrops commonly covered with white and orange lichen. Thickness approximately 200 m (650 ft).

Ch Hasmark and Cambrian (Cambrian)—Light gray to bluish gray, thin- to very thick-bedded, medium crystalline dolomite with minor magnesium limestone intervals. Basal part contains minor pebbles, succeeded upward by oolitic and psilotic intervals, argillite-marble, minor intraformational conglomerate, and thin quartz sandstone beds. Weathers very light gray to pale brown with a gritty, laminated surface. Minor replacement chert nodules and chert beds, a few centimeters to decimeters thick, distinguish this unit from Jefferson Formation (D). Exposed thickness is approximately 122 m (400 ft).

Ch *Silver Hill Formation* (Cambrian)—Thin-bedded, olive green to olive gray, micaceous shale and argillite with interbedded reddish quartz sandstone and siltstone. Mapped on the basis of float in northwest corner of the map, southwest of Humboldt Mountain. Thickness not determined but as much as 60 m (200 ft) in the southern Pioneer Mountains and locally absent as a result of erosion (Myers, 1952).

Cq Quartz sandstone (Cambrian)—Light gray to grayish orange, moderate to poorly sorted, medium- to coarse-grained quartzite with subordinate dark argillite and pale reddish quartzite interbeds. The gray quartzite has subangular to rounded quartz grains with rare (< 1–2 percent) dark grains, and 3–5 percent plagioclase feldspar. Commonly forms talus but where exposed, outcrop is medium- to thick-bedded, massive to cross-bedded, with localized lenses and discontinuous beds of well-sorted quartz granules and pebbles. Interbedded argillite is dark gray to grayish-green, laminated to ripple cross-laminated, with thin interbeds of fine- to coarse-grained, gray to pale reddish quartzite with up to 15 percent potillite feldspar, and rare quartzite and siltite pebble lenses that are often glauconitic. Argillite beds are poorly exposed but appear to be laterally discontinuous. Cambrian trace fossils, including Cruziana (trilobite feeding trail) and *Palaophycus striatus* (dovelling burrow), and several unidentified horizontal and vertical burrows, are relatively common, especially in float around and at the top of Humboldt Mountain (Fig. 3, A, B, C, D, E, F, G, H, I, J, K, L, M, N, O, P, Q, R, S, T, U, V, W, X, Y, Z). The argillite may correlate with the Silver Hill Formation and the quartzite with the Flathead Formation; however, the lower part of the sequence could be older. The upper and lower contacts are poorly exposed but estimated thickness is 244 m (800 ft).

C2YAR Argillite and quartzite (Cambrian and Mesoproterozoic)—Dark gray to grayish green argillite and shale interbedded with grayish red, pale reddish, and grayish green, fine- to coarse-grained, feldspathic and non-feldspathic quartzite. Argillite beds are typically laminated, 2–5 cm (2–2 in) thick, micaceous, and cap the quartzite beds or form intervals of fissile shale up to 3 m (10 ft) thick. Bedding surfaces can have nodular, wrinkled, or cracked structures interpreted as microbial mats (Kovalchuk, 2017). Quartzite beds range in thickness from 5 cm to about 0.3 m (2 in to 1 m). Common sedimentary structures include ripple marks, cross-bedding, and less commonly, distorted bedding and mudclips. Mostly mapped on basis of float but good exposures can be found in gullies and abandoned prospect pits in sections 7 and 18, T. 6 S., R. 10 W. Estimated thickness about 221 m (725 ft) based on a borehole (Pearson, 1996) in the west adjacent Ermon 7.5' quadrangle.

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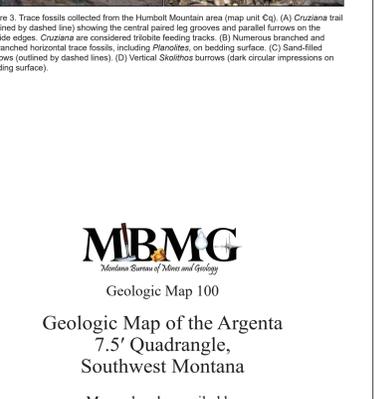
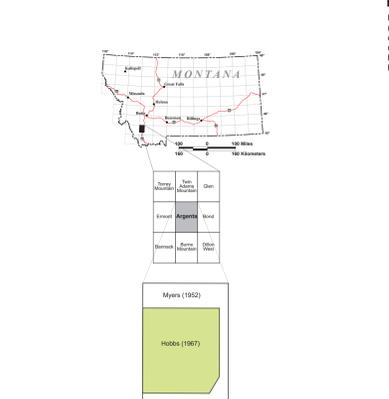
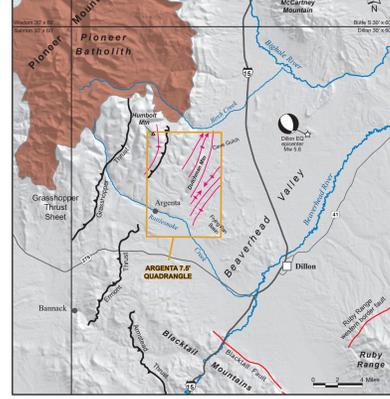
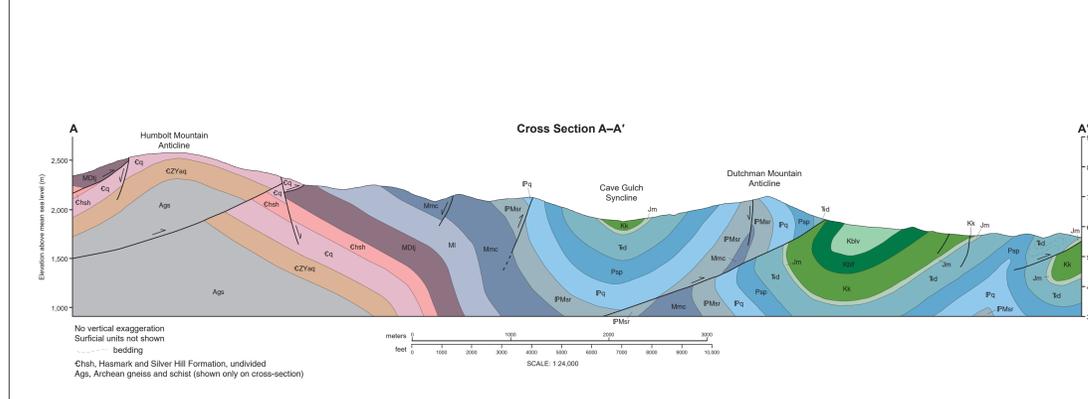


Figure 1. Cross section A-A' showing geological units and structures. The section shows the Humboldt Mountain Anticline, Cave Gulch Syncline, and Dutchman Mountain Anticline. Units are labeled with codes like Cq, C2YAR, Tr, etc. A scale bar and north arrow are included.

Figure 2. Simplified tectonic map of the area near the Argenta 7.5' quadrangle. Quaternary faults (Stockey and others, 2007) are shown in red.

Figure 3. Sources of mapping. The entire quadrangle was mapped by Ruppel and others (1993) for the Dillon 1:250,000-scale geologic map.

Figure 4. Trace fossils collected from the Humboldt Mountain area (map unit Cq). (A) Cruziana trail (outlined by dashed line) showing the central paired leg grooves and parallel furrows on the outside edges. (B) Argenta are considered trilobite feeding tracks. (C) Numerous branched and unbranched horizontal trace fossils, including Phylloporus, on bedding surface. (D) Sand-filled burrows (outlined by dashed line). (E) Vertical Skolithos burrows (dark circular impressions on bedding surface).