Beaverhead and Ravalli Counties, Montana



ene cene	QUATERNARY	
		> cenozoic
	>TERTIARY	
ene	J	J
	CRETACEOUS	} MESOZOIC
		> MESOPROTEROZOIC

Qaf Alluvial-fan and debris-flow deposits (Holocene to Late Pleistocene)—Angular to subrounded, poorly sorted, matrix-supported pebble to boulder gravel in

moderately sorted and stratified pebble to boulder sandy gravel above current stream level along the North Fork of the Salmon River, Pierce Creek, and the West Fork of Camp Creek. Upper surface is relatively smooth and flat. Locally grade upstream into *Qgo* deposits. Includes single high gravel deposit north of Cool Gulch that is older and possibly Tertiary in age. GLACIAL DEPOSITS

Outwash (Late Pleistocene)—Low areas of poorly sorted sand, gravel, cobbles, and boulders downstream from till deposits. Upper surface is steeper and less even than Qt deposits, but Qgo deposits grade into Qt deposits downstream so contacts between the two are approximate. Till deposits (Pleistocene)—Unsorted sand, gravel, cobbles, and boulders. Clasts are subangular to subrounded or faceted. Forms end and lateral

MASS-MOVEMENT DEPOSITS

Landslide deposits (Holocene to Pleistocene)—Unstratified, unsorted silty clay, gravelly silty clay, and boulders.

CHALLIS VOLCANIC GROUP Rocks of the Eocene Challis Volcanic Group formed from eruptions and to

a lesser extent sedimentation. They are widespread to the south where they were mapped by D.H. McIntyre, E.B. Ekren, and R.F. Hardyman in the Challis 1° x 2° quadrangle (Fisher and others, 1992). Here only remnants of originally more extensive volcanic units are present. Flow-banded rhyolite (Eocene)—Light tan to pinkish-gray flow-banded, largely

aphyric rhyolite (Figure 4). Crystal content generally sparse (less than 1 percent) but locally as much as 15 percent of quartz and potassium feldspar less than 1 mm in size. Trace amounts of biotite and hornblende present locally. Trace orange pumice as large as 2 mm, and lithics identical to matrix as large as 6 mm. Contains rare lithophysae. Lower two thirds of unit is pervasively flow-banded on a mm scale with banding highly distorted at some localities. Upper one third displays less distinct flow-banding. Weathers to form plates along welding surfaces. Presence of pumice and lithics,

tuff that flowed after emplacement. No dikes were observed cutting this unit. Sample 23DS19 from northwest of Lost Trail Pass was dated by U-Pb LA-MC-ICP-MS methods on zircon and returned an age of 49.3 ± 0.2 Ma Biotite dacite tuff (Eocene)—Light-gray to tan, moderately welded dacitic ash to lapilli tuff (Figure 5). Characterized by dark quartz crystals as large as 3

than 1 mm. Heterogeneous with varied proportions of crystals, lithics, and pumice. Crystals, which comprise 10 to 30 percent, are predominantly white plagioclase as large as 5 mm. Subordinate quartz is commonly dark and as large as 3 mm. Pseudohexagonal biotite and acicular hornblende, as much as 3 percent, are small (<1 mm). Rounded volcanic lithics as large as

of the rock. Rounded pumice as large as 1 cm is rare but locally comprises 20 percent of the unit. This tuff rests on an eroded surface of *Tdsr* and is cut by both rhyolite and dacite dikes. Mosolf and Kylander-Clark (2023b) reported a U-Pb zircon age of 52.2 \pm 0.4 Ma for a sample collected southwest of Saddle Mountain (sample AH21WW04). That sample contained 70.83 percent SiO₂ (Mosolf and others, 2023).

INTRUSIVE ROCKS

Quartz syenite (Eocene)—Single large dike in northwest part of map that was injected along, and later cut by, the Camp Creek fault. Termed the Saddle Mountain guartz syenite by Desmarais (1983) who noted that rock was composed of ~75 percent perthitic potassium feldspar, ~15 percent plagioclase, ~7 percent quartz, and lesser amounts of amphibole, biotite, pyroxene (aegirine/aegirine-augite), and opaque oxides that includes magnetite cubes as large as 0.5 mm. Texture ranges from coarse-grained equigranular to porphyritic with potassium feldspar phenocrysts. Textural and compositional variation indicates multiple injections of magma. Dated at 49.3 \pm 0.2 Ma (n = 38; MSWD = 0.8), using LA-MC-ICP-MS methods (sample AH21WW03 in Mosolf and Kylander-Clark, 2023b; see Symbols). SiO₂ content at dated locality is 64.08 percent (Mosolf and others, 2023). nate dacite dikes. Over 90 percent of this unit is rhyolite to rhyolite porphyry dikes, with minor dacite or dacite porphyry dikes and very rare Tgd country rock screens. See dike descriptions below for individual lithologies. *Tdsr* does not crop out as well as *Tds*, but where exposed dikes are tabular, less than 15 m (50 ft) thick, and are sub-vertical with northeast strikes. Along U.S. Highway 93 in Montana immediately north of the map boundary *Tdsr* is weathered to low, pinkish slopes that contrast sharply with the resistant dikes of *Tds* to the south. Date of 51.2 ± 0.4 Ma reported in Mosolf and Kylander-Clark, (2023b; sample AH21WW08) for a dike near the northern map boundary that contains embayed quartz phenocrysts 2 to 4 mm across and altered potassium feldspar(?) phenocrysts 5 to 8 mm in length. SiO₂ content of that sample is 67.01 percent (Mosolf and others, 2023). Mapped as a rhyolite to rhyodacite dike swarm by Desmarais (1983).

Tds Dike swarm (Eocene)—Mixed unit consisting of dacite dikes and small stocks, hyolite dikes, and minor amounts of country rock. Over 90 percent of this unit is dacite, dacite porphyry, rhyolite, and rhyolite porphyry dikes in roughly equal proportions; the remainder is country rock. See dike descriptions below for individual lithologies. The dike swarm is well exposed along U.S. Highway 93 near the northern map boundary. Dikes are parallel, tabular bodies as thick as 15 m (50 ft) that strike northeast and dip 70 degrees to the southeast. Cross-cutting relationships are rare. The exception is aphyric andesite dikes (Ta), only locally mapped, that appear to be late-stage intrusions. Similarity of dike composition, highly varied texture rarity of chilled margins, and similar dike geometry, are consistent with injection from multiple magma chambers, possibly with magma mixing, over a relatively short time span. Porphyritic dacite (Figure 6: 23RL043) from southwest of Lost Trail Pass was dated by U-Pb LA-MC-ICP-MS methods on zircon and returned an age of 52.7 ± 0.3 Ma (Figure 3B). This sample

Country rock within *Tds* is rare. North of Lost Trail Pass rare screens of *Tgd* or *Tmbg* separate dikes; south of the pass rare screens of *Yqcs*, *Tgd*, *Tmbg*, and *Kgdf* are within *Tds*. Dacite dikes grade into granodiorite (*Tgd*), which is locally mapped separately. Mapped in part as biotite granodiorite by Granodiorite (Eocene)—Fine- to medium-grained, equigranular to porphyritic biotite- and hornblende-biotite granodiorite. Grades into Tds where

Highway 93 dated at 52.4 \pm 0.2 Ma (n = 27; MSWD = 0.9), using LA-MC-ICP-MS methods (sample AH21WW05 in Mosolf and Kylander-Clark, 2023b; see Symbols). Sample contains embayed quartz phenocrysts 2 to 4 mm across and feldspar phenocrysts mostly 3 to 5 mm in length. SiO₂ content in that sample is 68.21 percent (Mosolf and others, 2023). That locality is slightly coarser grained and more equigranular than

Andesite dikes (Eocene)—Sparsely porphyritic andesitic dikes. Phenocrysts are plagioclase, biotite, and hornblende. Rhyolite dikes (Eocene)—Rhyolite dikes with sparse quartz and potassium feldspar phenocrysts. Magnetic susceptibility at one exposure is 0.19×10^{-3}

rhyolite dikes with conspicuous potassium feldspar phenocrysts. Phenocrysts that comprise as much as 60 percent of the rock are white plagioclase as large as 4 mm, lesser clear quartz as large as 5 mm, and pink euhedral potassium feldspar as large as 15 mm.

23RL030 45.6723 -113.9532 Foliated granodiorite

2RB406 45.6203 -113.8921 Mafic intrusive rocks

22RB402 45.6106 -113.7933 Muscovite-biotite granite

2LK008 45.5365 -113.8280 Mafic intrusive rocks?

Dacite dikes (Eocene)—Dacite dikes with sparse plagioclase, quartz, ornblende, and biotite phenocrysts. Proportions of phenocrysts vary. Texture spans from nearly aphyric to nearly phaneritic, which closely resembles Tgd. Magnetic susceptibility at two exposures is 0.18×10^{-3} SI. Tdp Porphyritic dacite dikes (Eocene)—Highly porphyritic dacite dikes with plagioclase, guartz, hornblende, and biotite phenocrysts. Proportions of phenocrysts vary. Textures range from highly porphyritic to nearly equigranular, which closely resembles that of Tgd. Magnetic susceptibility at two exposures is 0.21×10^{-3} SI.

Tmbg Muscovite-biotite granite (Paleocene)—Massive muscovite-biotite granite to granodiorite. Locally sparsely porphyritic with potassium feldspar phenocrysts. Minor to trace amounts of muscovite. Pegmatite and leucocratic granite are present locally. Desmarais (1983) mapped this as the Trail Creek pluton, which is more extensive east of this map. He reported weakly zoned plagioclase (An28-30) and common myrmekite. SiO₂ content of sample collected 7 km (4 mi) to the east-southeast in the Big Hole Pass quadrangle is 71.6 percent (sample 22RB402, Table 2). Magnetic susceptibility of two exposures is 0.21 and 0.48 \times 10⁻³ SI. Locality along State Highway 43 12 km (7 mi) east of the quadrangle dated by U-Pb methods at 63.3 ± 0.6 Ma (sample CE20BHB15 in Mosolf and Kylander-Clark, 2023a). Kgdf Foliated granodiorite (Cretaceous)—Foliated to massive biotite granodiorite Figure 7). Anhedral and relatively large biotite (1-4 mm). SiO₂ content is

69.7 percent (23RL030, Table 2). That sample from exposure on U.S. Highway 93 southwest of Lost Trail Pass was dated by U-Pb LA-MC-ICP-MS methods on zircon and returned an age of 76.8 \pm 0.4 Ma (Figure 3C). Magnetic susceptibility at that locality is 1.21×10^{-3} SI. Unit is older and more mafic than *Tmbg* and possibly correlative with the Martin Creek pluton of Desmarais (1983) that has been dated at about 78 Ma (but with a large error) on two samples. Kcs Calc-silicate veins (Cretaceous?)—Red-brown weathering coarse-grained actinolite-diopside-calcite veins or dikes. Contacts sharp and linear (Figure

8). Median magnetic susceptibility measured at four exposures (one was three hand samples) is 8.77×10^{-3} SI. Known only in Yg in southwest corner of map and south-adjacent Gibbonsville quadrangle (Stewart and others, in prep). Possibly derived from remobilization of carbonate within structurally or stratigraphically underlying Yqcs. Ymi Mafic intrusive rocks (Mesoproterozoic)—Fine- to medium-grained mafic

likes in southeast part of quadrangle. White plagioclase as much as 30 percent. Original hornblende and pyroxene partly altered to amphibole, chlorite, and epidote, likely during metamorphism. Western dike appears to have intruded along and subsequently been cut by the Anderson Creek fault. Sample 22RB406 collected near the state line 530 m (1,700 ft) south of the quadrangle was dated by U-Pb LA-MC-ICP-MS methods on zircon. Zircon yield from this sample was poor. Of nine grains analyzed, seven yielded similar ages with a 207 Pb/ 206 Pb weighted mean of 1264 ± 15 Ma (Figure 3D). One concordant (likely inherited) grain yielded a ~1747 Ma age and one discordant grain showed Cretaceous or younger Pb-loss. The same sample contained 52.7 percent SiO₂. A sample from a smaller mafic dike of uncertain age collected 10 km (6 mi) to the southeast of the quadrangle contained 50.8 percent SiO₂ (sample 22LK008, Table 2).

METASEDIMENTARY ROCKS

Mesoproterozoic metasedimentary rocks comprise most of the bedrock in the southern part of the quadrangle. They were metamorphosed to biotite grade but vary in grain size and biotite content. Unit names follow the stratigraphic nomenclature (Burmester and others, 2016b) that uses the coarsest unit, the Swauger Formation, to distinguish otherwise similar rocks that are present above and below it.

wauger Formation (Mesoproterozoic)-Light-gray to white, fine- to medium-grained, poorly-sorted quartzite with more feldspar, as much as 50 percent, than typical for this unit; potassium feldspar typically greater than plagioclase. Where mapped as Swauger Formation undivided (Ys of Stewart and others, 2014) quartzite beds generally fine upward within laterally defined by heavy mineral laminae with hematite. Approximately 3,400 m (11,000 ft) thick in the Allan Mountain quadrangle to the south (Stewart and others, 2014).

Yg **Gunsight Formation (Mesoproterozoic)**—Quartzite, siltite, and minor argillite. Mapped as Gunsight Formation by Stewart and others (2014) and described as gray, gray-green, and tan, well sorted, fine-grained feldspar-rich quartzite. Features include abundant heavy mineral laminations and soft-sediment deformation. The quartzite is well sorted, and is laminated to thinly bedded, commonly showing ripples, climbing ripples, load casts, and trough and planar cross-bedding. Median magnetic susceptibility measured at seven exposures and ten hand samples is 0.10×10^{-3} SI. Upper contact gradational above approximately 400 m (1,300 ft) of green to light gray, siltite to fine-grained guartzite (Stewart and others, 2014). Thickness on the map uncertain but estimated to be approximately 4,500 m (15,000 ft) (Stewart and others, 2014). Stratigraphic position under the Swauger Formation supports assigning this unit to the highest unit of the Lemhi

> Quartzite and calc-silicate rocks (Mesoproterozoic)—Quartzite, siltite, calc-silicate rocks, and minor schist and phyllite. Bedding on cm to dm scale (Figure 9). Decreasing abundance of calc-silicate rocks northward, interpreted to be stratigraphically downward. Includes phyllite, schist, and gneiss in the northern and eastern exposures near contacts with plutonic rocks. Calc-silicate rocks are resistant to weathering and commonly crop out. They contain abundant actinolite along with plagioclase and local scapolite. Median magnetic susceptibility measured at eight exposures and two hand samples is 0.25×10^{-3} SI. Susceptibility was higher for two exposures within Tds: 4.18×10^{-3} SI. Thickness unknown as no lower or upper stratigraphic contact has been found, but a minimum of 1,000 m (3,280 ft) is likely. Previously mapped as calc-silicate, quartzite, and siltite of Dahlonega Creek (Stewart and others, 2014). If stratigraphically below Yg, lithologies suggest correlation with the Yellow Lake Formation, which contains carbonate, and possibly parts of the Big Creek Formation, largely very fine grained quartzite and siltite, both of the Lemhi Group (Burmester and others, 2016b). A similar stratigraphic level within strata mapped as the Yellowjacket Formation also contains carbonate (Ekren, 1988). Alternatively, could be below those units if normal displacement on the Cool Gulch fault was very large, perhaps the Piegan Group, mapped 14 km (9 mi) and 35 km (22 mi) northeast of this map (Elliott and Lonn, in prep.). Another possible correlative if the Cool Gulch and Anderson Creek faults had an earlier thrust history is the quartzite, argillite, and minor calc-silicate mapped above the Swauger Formation in the West Pioneer Mountains to the east (Lonn and Scarberry, 2022). However, that unit has medium quartz grains typical of the Lawson Creek Formation (Hobbs, 1980) that are absent

STRUCTURE

This map and the Gibbonsville quadrangle to the south (Stewart and others, 2014; Stewart and others, in prep) share Ys, Yg, Yqcs, Ymi, Kcs, and the Anderson Creek fault. The Cool Gulch fault and the Lick Creek fault (on the Gibbonsville map) are the southern contacts of Yqcs with Yg or Ys. Understanding the local structure therefore depends on both maps. The general map pattern of Mesoproterozoic units west of the Anderson Creek fault is of a southward dipping homocline, with Yqcs under Yg under Ys. In the few outcrops where we were able to determine top directions, all strata were south to southwest facing. Although *Yg* is separated from *Yqcs* by the Cool Gulch fault, *Yqcs* could be stratigraphically below *Yg* if the fault were minor and only omitted part of a homoclinal section.

TWIN CREEK FAULT The Twin Creek normal fault in the southwest part of the map can be traced

14 km (9 mi) south through the Gibbonsville quadrangle where it was first mapped (Stewart and others, 2014; Stewart and others, in prep). In the Gibbonsville quadrangle it places Challis Volcanic Group strata in the east against Mesoproterozoic metasedimentary rocks in the west. Brecciated fault rock at the mouth of Cool Gulch is attributed to this structure. The northern extent beyond Cool Gulch is uncertain.

COOL GULCH FAULT The Cool Gulch fault, formerly the northern offset western segment of the

Lick Creek fault (Stewart and others, 2014), is the contact between Yg and *Yqcs* in the southwestern part of the map. Here the Cool Gulch and Lick Creek faults are given separate names because of uncertainty that they are the same structure. Because of the uncertainty in the stratigraphic position of the Yqcs unit, several options exist for the nature of this contact. If Yqcs is older than Yg (middle to lower Lemhi Group) then it is likely a minor down-to-the-south normal fault. If *Yqcs* is younger than *Yg* and *Ys*, and equivalent to the Lawson Creek and Apple Creek formations, then it is a major structure. An exposure west of the North Fork Salmon River contains 50° south-southwest plunging structures that look like tight folds with amplitudes on the cm scale (Figure 10) in steeply dipping, thinly bedded quartzite with thinner schistose tops. Whether this deformation records faulting or folding is unclear. To the east the fault is offset by the Twin Creek fault and other normal faults so appears older than those, the intrusive rocks to the northwest, and Td that intrudes it east of the Twin Creek fault. It appears to cut out Yg eastward. The fault terminates at the Anderson Creek fault, which must have offset it or accommodated its displacement. Because of the apparent antiquity of the Anderson Creek fault (see below), the Cool Gulch fault may also be an old (Mesoproterozoic) structure.

ANDERSON CREEK FAULT

The Anderson Creek fault is mapped across the southeast corner of the map south into the Gibbonsville quadrangle (Stewart and others, 2014) and the North Fork quadrangle (Lonn and others, 2013). The northern extent of the fault is uncertain. We show it being truncated by the Pierce Creek fault, but it may instead be intruded out by the *Tmbg* unit. The mafic intrusion that appears to have intruded along the fault has been provisionally dated at ~1,260-1,270 Ma (see Ymi description above). If the date is correct, this structure had a long history. Both dextral and down-on-the-west normal slip was proposed by Stewart and others (2014) based on stratigraphic arguments and 11 km (7 mi) offset between the Lick Creek fault contact of *Ys* with *Yqcs* in the Gibbonsville quadrangle and the Cool Gulch fault contact of Yg with Yqcs near the south edge of this map. The dissimilarity of those contacts makes them a poor piercing point, however. Stewart and others (2014) also cited Bacorn (1905) and Mayerle and Close (1993) for evidence from underground mines in the Gibbonsville area that indicated right-lateral strike-slip motion. Post-Eocene dip-slip motion is supported by the presence of Challis Volcanic Group strata west of the fault in the Gibbonsville and North Fork quadrangles (Stewart and others, 2014; Lonn and others, 2013). Net slip depends on the stratigraphic position of Yqcs versus Yg.

CAMP CREEK FAULT The Camp Creek normal fault strikes southwest and dips northwest across

the northwest part of the quadrangle. Along part of its trace, it is the northwest contact of the quartz syenite unit (Tqs). Northwest of it the Eocene dike swarm is dominated by rhyolite (*Tdsr*), as opposed to the dacite-dominated dike swarm to the southeast (*Tds*). It appears to mark a transition from more north-trending faults north of this quadrangle to a northeast structural grain crossing the central part of the Lost Trail Pass quadrangle.

MINERALIZATION

Several prospects are present in the quadrangle, but no significant amount of mining has taken place. Unpublished maps and documents for mining properties are available by examining "Property Details" through the nteractive map ("Mines" web app) on the Idaho Geological Survey website (https://www.idahogeology.org/webmap). Property code (e.g., DI0001) is given below to assist with website searches.

KOPER KYUTE PROSPECT (DI0001)

The Koper Kyute prospect in the south-central part of the quadrangle is developed by an upper trench and a lower pit that exposes a northwest-southeast zone of probable calc-silicate rock with local copper stain. Exposures in the lower pit were initially thought to be dike, but thin section examination indicates that hornblende, quartz, and epidote are the primary minerals and titanite, plagioclase, pyroxene, and calcite are minor constituents. Rock has varied and high magnetite content. Median magnetic susceptibility from 15 measurements (Yqcs-KK, Table 1) is 16.19×10^{-3} SI; the highest reading of $1,732 \times 10^{-3}$ SI shows that about half of that rock measured is magnetite (reference data from Hunt and others, 1995).

Possibly a magnetite skarn zone. One of two samples from the lower pit has an anomalous Cu concentration (1,716 ppm; 23RL028a, Table 2). Although no intrusive rocks are present at the lower pit, minor felsic intrusive float was found in the upper trench and there could be additional intrusions at depth related to the nearby *Td* and *Tr* dikes and the dike swarm (*Tds*) to the northwest.

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Kgdf Lost Trail Pass 69.70 0.359 15.35 2.70 0.060 0.99 3.34 3.89 2.29 0.109 98.79 1.08 6 13 4 38 1043 69 512 169 12 10.8 18 11 54 17 45 83 11 27

Ymi? Big Hole Pass 50.77 1.039 12.77 10.07 0.171 10.05 9.98 1.44 1.05 0.088 97.44 2.15 218 671 33 274 246 42 126 88 21 4.2 16 90 76 7 11 25 2 14

 Tmbg
 Big Hole Pass
 71.58
 0.242
 14.87
 1.98
 0.084
 0.59
 2.10
 3.86
 3.27
 0.094
 98.67
 0.88
 5
 7
 4
 26
 763
 107
 271
 102
 22
 15.7
 17
 1
 45
 24
 24
 41
 7
 16
 0

52.67 1.474 12.75 10.42 0.149 8.45 9.38 2.20 1.21 0.152 98.85 1.05 170 436 28 285 214 46 256 120 20 7.6 18 75 66 1 14 33 2 19







Gray to pinkish-gray highly porphyritic us potassium feldspar phenocrysts. h as 60 percent of the rock are white clear quartz as large as 5 mm, and pink e as 15 mm.	Gibbonsville and No and others, 2013). N versus Yg. The Pierce Creek fat southern part of the o we have tentatively s	the presence of Challis Volcanic Group strata west of the fault in the Gibbonsville and North Fork quadrangles (Stewart and others, 2014; Lonn and others, 2013). Net slip depends on the stratigraphic position of <i>Yqcs</i> versus <i>Yg</i> . PIERCE CREEK FAULT The Pierce Creek fault is mapped as a northeast-striking structure in the southern part of the quadrangle. Its northeastern extension is uncertain but we have tentatively shown it to be the northern terminus of the Anderson Creek fault. Its northeastern trace is based on field relations east of the map.					ti s C F	Figure 1. Pre-Mesozoic bedrock geology around Salmon, Idaho. The coarsest clastic unit (Ys = Yh) separates lower formations of the Lemhi Group (in blue) from higher Lawson Creek and Apple Creek formations (greens and yellows). Units are combined where scale or previous mapping makes separating them impractical. Geographic locations are in italics: AM—Allan Mountain; BC—Big Creek; CL—Cowbone Lake; GM—Goat Mountain; GP—Gunsight Peak; HC—Hayden Creek, JL—Jahnke Lake; LC—Lawson Creek; LM—Lake Mountain; LP—Lem Peak; MC—Moose Creek; MM—Mogg Mountain; RM—Ramsey Mountain; WFB—West Fork Bitterroot River; YC—Yearian Creek; YL—Yellow Lake. Dashed rectangles outline the Salmon 1° x 2° (orange) and Lost Trail Pass (blue) maps. Modified after Burmester and others (2023) to include insights from field work in 2023.								22RI 23RI 1: 23RI 22R 23RI 23RI	2DS72 Bms004b Bms0451 2DS71 Bms0548 Bms006 Bms0455 Bms0449 Bms005		Td Yqcs Yqcs Yqcs Yqcs Yqcs Yqcs Yqcs Yqcs	45.6511 45.6555 45.6648 45.6520 45.6479 45.6448 45.6632 45.6669 45.6485	-113.9503 -113.9714 -113.9383 -113.9509 -113.9846 -113.9706 -113.9373 -113.9365 -113.9704						
Table 2. Major Sample Latitude Longitude 23RL028a 45.6635 -113.9375 Ouartzite	oxide and trace-element ana Unit name Map unit	, 1		the Lost Tr. 02 Al2O3			5	,		, I		ed by X-ra	,	escence	method	5 at Wash Ba	ington Sta	te Univer Zr	rsity. Coord	linate datur Ga Cu		Се	Th	Nd U			

Measurement ID

23RBms0448

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12MDM068

12MDM070

12DS74

23RBms0535

12DS69

12RL238

12DS73

23RBms0551

12MDM066

23RBms0553g

23RBms0549

12RI 239

23RBms0454

23RBms0459

22RBms003

23RBms0517

23RBms0514

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