

GEOLOGIC MAP OF THE BOZEMAN 30' x 60' QUADRANGLE
SOUTHWESTERN MONTANA

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Compiled and mapped by

Susan M. Vuke¹, Jeffrey D. Lonn¹, Richard B. Berg¹ and Christopher J. Schmidt²

¹Montana Bureau of Mines and Geology ²Western Michigan University

GEOLOGIC SOURCE MAPS AND INDEX OF 7.5' QUADRANGLES
BOZEMAN 30' x 60' QUADRANGLE

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	45°30'									

Numbers above correspond with reference list on next page.

GEOLOGIC MAP SOURCES BOZEMAN 30' x 60' QUADRANGLE

Numbers below correspond with index of map sources on previous page.

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Entire quadrangle

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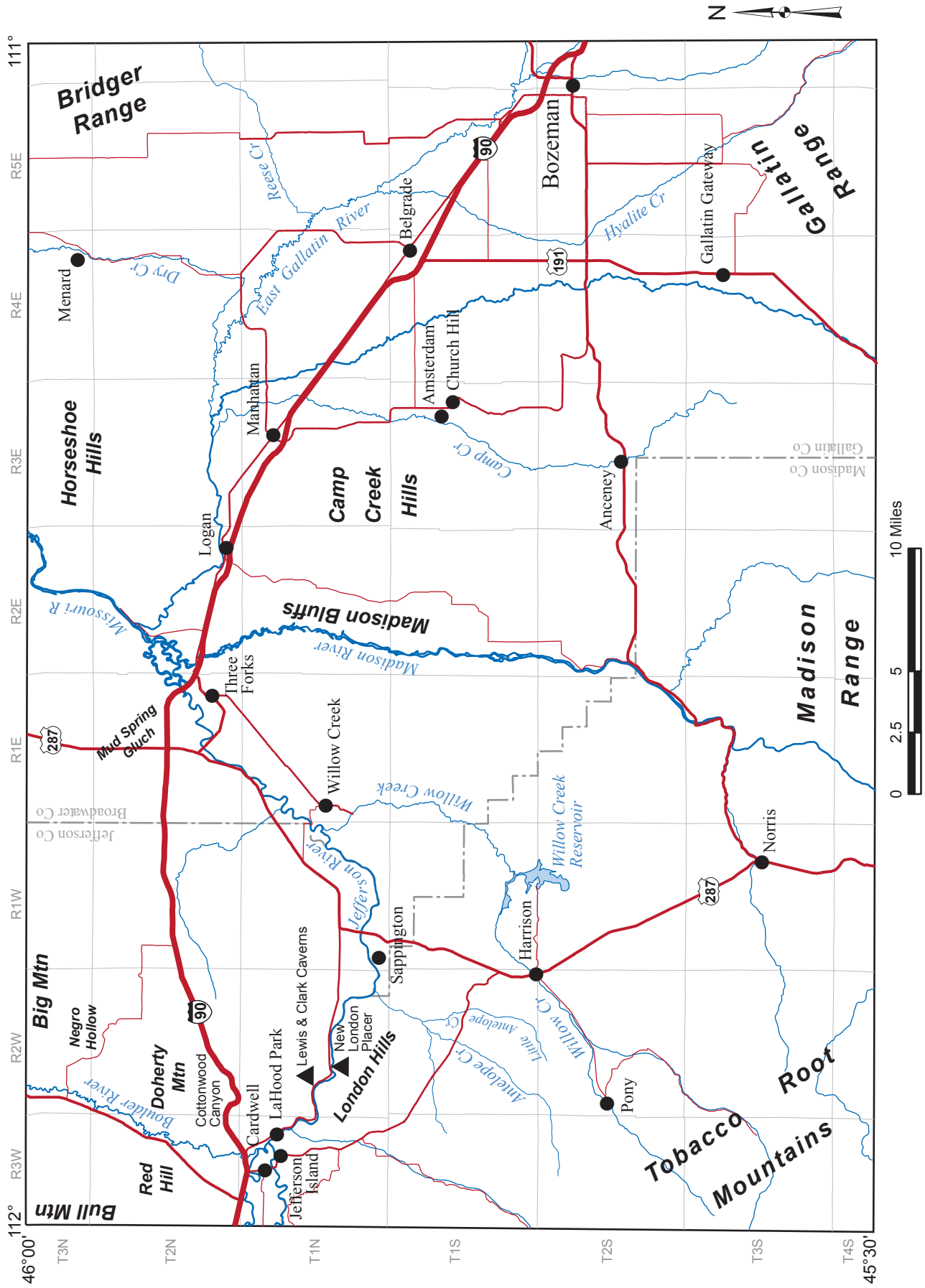


Figure 1. Location map for Bozeman 30' x 60' quadrangle.

DESCRIPTION OF MAP UNITS
BOZEMAN 30' x 60' QUADRANGLE

- Qal Alluvium (Holocene)**—Light gray to light brown gravel, sand, silt, and clay deposited in stream and river channels, on their floodplains, and on low terraces as much as about 6 m (20 ft) above modern streams and rivers. Moderately sorted to well sorted. Larger clasts subangular to well rounded. Composition varies, but includes clasts of Archean metamorphic rocks, Precambrian orthoquartzite, Paleozoic limestone and quartzite, vein quartz, and volcanic rocks. Clasts of some small streams originating in Tertiary uplands are dominantly granule size and smaller, and may include rip-up clasts.
River alluvium:
Gallatin and East Gallatin Rivers: Rounded to well-rounded small boulders, cobbles, gravel, sand, silt, and clay, dominantly composed of Archean metamorphic rocks, and dark-colored volcanic rocks, with subordinate Paleozoic limestone, and Precambrian Belt rocks.
Madison River: Subrounded to rounded, gravel and sand with clasts rarely larger than cobble size (Robinson, 1963) dominantly composed of Archean metamorphic rocks, dark-colored igneous rocks, Paleozoic limestone, quartz, and chert.
Tributaries of Gallatin and East Gallatin River: Clast composition varies. Clasts of tributaries from the Tertiary of the Camp Creek Hills are derived primarily from the pebbles and cobbles of Tertiary fluvial deposits. Dry Creek and its tributaries contain clasts derived from the Maudlow Basin, Horseshoe Hills, Bridger Range, and local Tertiary and older Quaternary deposits. Estimated thickness of Holocene alluvium of Gallatin River is 15-25 m (50-80 ft) based on thicknesses in well logs where Qal rests directly on bedrock.
Jefferson River: Dominantly subrounded to rounded and fairly well sorted with clasts rarely larger than cobble size. Thickness probably less than 15 m (50 ft) in most areas.
- Qc Colluvium (Holocene)**—Unstratified, unconsolidated, poorly sorted, angular to subangular clasts derived from local sources. Thickness unknown, but maximum thickness is probably less than 10 m (33 ft).
- Qls Landslide deposit (Holocene)**—Unstratified, unsorted mixtures of sediment that moved downslope through mass wasting processes. Includes rotated or slumped blocks of bedrock and surficial sediment, earthflow deposits, and mudflow deposits. Color and lithology and grain size reflect that of parent rock and transported surficial material. Thickness probably less than 30 m (100 ft).
- Qpa Paludal deposit (Holocene)**—Sand, silt, and organic matter deposited in swamp, pond, or small lake environment. Thickness probably less than 10 m (33 ft).
- Qdf Debris-flow deposit (Holocene)**—Angular, subangular and subrounded clasts of poorly sorted, locally derived boulders and cobbles in a matrix of fine-grained sediment. Thickness probably about 15 m (50 ft) in thickest part.
- Qe Eolian deposit (Holocene and Pleistocene)**—Yellowish gray to very pale orange, silt and clay-size sediment with scattered grains and thin lenses of rounded fine-grained sand. Dominantly volcanic glass, quartz, and clay minerals with minor amounts of mica, feldspar, and calcite, also silt, and clay. Thinly mantles many more areas than shown on

map. Thickness highly variable. Locally as much as 30 m (100 ft), but generally 1 m (3 ft) or less.

- Qac Alluvium and colluvium, undivided (Holocene and Pleistocene)**—Locally derived sediment on slopes. Color reflects that of parent material. Ranges from clay and silt to gravel, depending on source. Thickness generally less than 6 m (20 ft).
- Qgr Gravel deposit (Holocene and Pleistocene)**—Variable deposits of gravel, sand, silt, and clay that include alluvium, pediment veneer, colluvium, outwash, and fan deposits. Clast composition, rounding, and sorting vary depending on deposit location and type. The most extensive deposits are in the Gallatin Valley and include dissected blankets of sub-rounded to rounded pebbles, cobbles, small boulders, and rare large boulders. Composition is dominantly orthoquartzite, vein quartz, quartz-rich gneiss, and dark volcanic rocks. North of the Jefferson River, most gravel deposits are fan and alluvial deposits that overlie pediment surfaces. Clast composition reflects local sources and clasts range from angular to rounded depending on the type of deposit. Near the headwaters of the Missouri River, some of the deposits on pediment surfaces consist dominantly of angular limestone clasts. Thickness ranges from pediment veneer about 2.5 cm (1 inch) thick to valley fill as much as 120 m (400 ft) thick.
- Qaf Alluvial fan deposit (Holocene and Pleistocene)**—Deposit with preserved fan morphology at break in slope, composed of a heterogeneous mixture of subangular to moderately rounded coarse clasts as large as boulders, and fine-grained sediment (sand, silt, and clay) that is generally concentrated near fan margins. Clasts derived from adjacent uplands. Estimated thickness is about 30 m (100 ft) at thickest part.
- Qab Alluvium of braid plain (Holocene and/or Pleistocene)**—Rounded to well-rounded, dominantly cobble gravel with clasts as large as boulders, and sand, silt, and clay; mostly composed of Archean metamorphic rock fragments, and dark-colored volcanic rocks, with subordinate Paleozoic limestone and Proterozoic Belt rocks. Clast lithologies in general order of decreasing abundance include Precambrian metamorphic rocks, mafic volcanic rocks, dacite(?) porphyry, quartzite, sandstone, limestone, and chert. Two wells in this unit adjacent to the Gallatin River indicate thicknesses of Quaternary alluvium overlying Tertiary deposits of 9.5 m (31 ft) and 65 m (215 ft) (Hackett and others, 1960).
- Qat Alluvial terrace deposit (Holocene and/or Pleistocene)**—Subrounded and rounded pebbles, cobbles, and sand with some thin beds of clayey silt. Underlies distinct terraces adjacent to and at elevations higher than modern streams. Clast lithologies variable, but may include Precambrian metamorphic rock, igneous rock, quartzite, and sandstone. Thickness about 6 m (20 ft).
- Qafh Hyalite Alluvial Fan (Pleistocene)**—Deposit with distinct fan morphology composed of light gray to light brown gravel, sand, silt, and clay deposited where Hyalite Creek crosses an abrupt change in slope gradient (where the fault-bounded northern Gallatin Range meets the Bozeman Valley) extending for about 11 km (7 mi) into the valley. Distribution of clast sizes varies. In general, coarse-grained sediment is dominant near the head of the fan and fine-grained sediment near the margins. Clasts are dominantly matrix supported and poorly sorted, although sediment deposited in distributary channels is moderately to well sorted and clast supported. Large clasts subrounded to rounded. May be as much as 60 m (200 ft) thick.

- Qafo** **Alluvial-fan deposit, older than Qaf (Pleistocene)**—Light brown, gray, and locally reddish gray, angular and subangular, locally derived gravel in a coarse sand and granule matrix. Clast size ranges from pebble to small boulder. Fan morphology dissected. Maximum thickness probably about 45 m (150 ft).
- Qabo** **Braid plain alluvium, older than Qab (Pleistocene)**—Rounded to well-rounded, dominantly cobble gravel with clasts as large as boulders, and sand, silt, and clay; mostly composed of clasts of Archean metamorphic rock, and dark-colored volcanic rock, with subordinate Paleozoic limestone and Proterozoic Belt rocks. Clast lithologies in general order of decreasing abundance include Precambrian metamorphic rocks, mafic volcanic rocks, dacite(?) porphyry, quartzite, sandstone, limestone, and chert. A well in this unit indicates a thickness of 9 m (30 ft) of alluvium overlying Tertiary deposits.
- Qalo** **Alluvium, older than Qal (Pleistocene)**—Subrounded to rounded, fairly well sorted gravel with relatively few clasts larger than cobble size, and sand. Deposit covered by organic material, silt, and mud in some areas. Late Pleistocene bison and Holocene animal bones have been found in these deposits near Three Forks (Robinson, 1963). Thickness probably no more than 15 m (50 ft) in the Three Forks area (Robinson, 1963).
- Qlso** **Landslide deposit, older than Qls (Pleistocene)**—Unstratified, unsorted mixtures of yellowish brown Tertiary sediment that moved downslope through mass wasting processes north and east of Bozeman. Offsets from landslide displacement apparent in roadcut exposures. Thickness probably less than 30 m (100 ft).
- Qep** **Eolian and pediment deposits, undivided (Pleistocene)**—Yellowish gray, fine-grained tuffaceous sand and silt overlying pediment veneer of coarser sediment. Thickness generally less than 1 m (3 ft).
- Qm** **Mantle (Pleistocene)**—Regolith and lag deposits with clasts as large as boulders derived from underlying Tertiary deposits and from Quaternary debris-flow deposits, subordinate water-transported deposits, and colluvium.
- Qp** **Pediment deposit (Pleistocene)**—Sediment veneer that ranges from angular to rounded, dominantly pebble size or smaller clasts, to gravel on pediment surface.
- Qg** **Glacial deposits, undivided (Pleistocene)**—Poorly sorted, angular to rounded, unconsolidated clasts of dominantly cobbles and boulders, but also pebbles, sand, silt, and clay. Large clasts are of dominantly Archean metamorphic rocks, and igneous rocks of the Tobacco Root Batholith. Includes unstratified till and stratified outwash, and thin cirque lake deposits that probably include some Holocene sediment. Thickness of deposits less than 45 m (150 ft).
- Qgk** **Glacial kame deposit (Pleistocene)**—Moderately well sorted and stratified deposit in Tobacco Root Mountains associated with other glacial features.
- Qgt** **Glacial till (Pleistocene)**—Unconsolidated, poorly sorted, angular to rounded clasts of dominantly cobbles and boulders, but also pebbles, sand, silt, and clay. Large clasts dominantly composed of Archean metamorphic rocks, and igneous rocks of the Tobacco Root Batholith. Includes some stratified outwash. Thickness less than 45 m (150 ft).

- Qgo** **Glacial outwash deposit (Pleistocene)**—Moderately to well sorted, subrounded to well rounded gravel immediately downslope from glacial till (Qgt deposits).
- Qrg** **Rock glacier deposit (Pleistocene)**—Angular rock debris cemented by interstitial ice.
- QTaf** **Alluvial fan deposit (Pleistocene and/or Pliocene)**—Deposit similar to Qaf, but underlies Qaf, and is separated from it by a pediment surface. Thickness as much as 35 m (115 ft).
- QTafy** **Alluvial fan deposit, younger than QTafo (Pleistocene and/or Pliocene)**—Overlies QTafo in Negro Hollow area. Light gray unconsolidated, angular to subangular, and subordinate subrounded gravel. Clasts are almost exclusively locally derived from Paleozoic limestone but also include locally derived chert and Quadrant Formation. Thickness probably less than 30 m (100 ft).
- QTafo** **Alluvial fan deposit, older than QTafy (Pleistocene and/or Pliocene)**—Underlies QTafy in Negro Hollow area. Light gray unconsolidated, angular to subangular, and subordinate subrounded gravel, and well cemented breccia, especially at the base. Basal breccia matrix is reddish brown in most places and patches of reddish brown matrix also occur locally, higher in the unit. Clasts are almost exclusively locally derived Paleozoic limestone but also include locally derived chert and Quadrant Formation. Includes slightly bentonitic, pinkish gray and very light gray, tuffaceous sandstone west of the southern part of Big Mountain. Thickness probably less than 30 m (100 ft).
- QTat** **Alluvial terrace deposit (Pleistocene and/or Pliocene)**—Angular clasts of Elkhorn Mountains Volcanics associated with a pebble gravel in a very limited linear distribution on the tops of low hills along a tributary of Spring Gulch in the northwestern part of the map. A Miocene Barstovian fossil was found in this unit (Robinson, 1963). Thickness about 1 m (4 ft).
- QTep** **Eolian, paleosol, and pediment deposits (Pleistocene and/or Pliocene)**—Dark brown, clayey silt and fine-grained eolian sand with sparsely distributed, matrix-supported pebbles and some cobbles that alternates with weakly developed paleosols. Pediment surface on deposit veneered with pebble-size and smaller clasts transported by sheetwash. Thickness probably less than 15 m (50 ft).
- QTge** **Gravel and eolian deposits, undivided (Pleistocene and/or Pliocene)**—Deposits in Jefferson Canyon with two distinct components – a gravel with mostly boulders, overlain by fine-grained sediment of probable eolian origin. At the New London placer, near Lewis and Clark Caverns, clast composition includes Archean metamorphic, Belt Supergroup quartzite, subangular and subrounded clasts of LaHood Formation conglomerate, individual quartzite clasts recycled from the LaHood Formation conglomerate, and Paleozoic limestone.
- QTdf** **Debris-flow deposit (Pleistocene and/or Pliocene)**—Angular, subangular, and subrounded clasts of poorly sorted, locally derived boulders and cobbles in a matrix of fine-grained sediment.
- QTgr** **Gravel deposits (Pleistocene and/or Pliocene, and patchy remnants of Miocene)**—Near Willow Creek, gravel consists of angular to rounded matrix-supported cobbles, most of which are white quartz, schist, and amphibolite. Some of the clasts have caliche rinds. Matrix is sand and silt. An immature, matrix-supported breccia with angular to

subangular granules, pebbles, and cobbles of Archean metamorphic rocks is below the gravel and included in the map unit. Combined thickness of both deposits is 6-12 m (20-40 ft). The deposit in the Tobacco Root Mountains contains matrix-supported, well-rounded pebbles, cobbles, and boulders up to 9 m (30 ft) thick in a sandy matrix. Clast composition includes Archean metamorphic rocks, granodiorite of the Tobacco Root Batholith, and well-cemented sandstone. Deposits west of Gallatin Gateway have the same clast composition as pebbles and cobbles in the underlying Tertiary rocks, and may be local deposits related to faulting. In addition, several fossils indicate that there are patchy remnants of Miocene deposits (Robinson, 1963) that are included in this map unit west of the Madison River. Thickness about 9 m (30 ft).

Tertiary deposits of unknown age

Ts Sediment or sedimentary rocks, undivided (Tertiary)

Trs Rhyolite sediment (Tertiary)—Very light gray to white, well-bedded, well-indurated siltite, sandstone, and clast-supported sedimentary breccia composed almost entirely of rhyolite vitrophyre; clasts are angular and as much as 0.9 m (3 ft) long. Interlayered and locally intermixed with abundant white air-fall ash. Exposed thickness about 30 m (100 ft).

Trb Red Bluff Formation (Tertiary) (Kellogg, 1994):

Upper: White and light yellowish brown siltstone, sandstone, conglomerate, and subordinate, but conspicuous, brick red and maroon mudstone and siltstone; locally tuffaceous. Most clasts are Archean gneiss, and vein or pegmatite quartz, with some quartz monzonite and granodiorite from Tobacco Root Batholith. Locally, highly silicified.

Lower: Subrounded to well-rounded, matrix supported, bouldery diamictite. Clast composition dominantly quartz monzonite or granodiorite of the Tobacco Root Batholith, with matrix of immature decomposed quartz monzonite or granodiorite. Boulders are very large, some as much as 15 m (50 ft) wide. Type locality is within quadrangle in W¹/₂ sec. 13, T. 3 N., R. 1 W.

Sixmile Creek Formation (younger than Hemingfordian unconformity)

Tscr Reese Creek member (Miocene: Clarendonian and Hemphillian)—Basal cobble or small boulder clast-supported conglomerate or gravel cemented by calcium carbonate, or unconsolidated with clasts coated by calcium carbonate overlain by orangish tan tuffaceous mudstone with lenses of fine-grained sandstone. Fine-grained upper part of unit best exposed near the mouth of Reese Creek. In most other areas, the poorly resistant, fine-grained part is poorly exposed. A vitric ash was sampled from near the mouth of Reese Creek that provided a K/Ar date of 8.9±0.4 Ma (youngest Clarendonian or oldest Hemphillian). The ash has an abrupt basal contact and climbing ripples at the top, and is 0.9-1.5 m (3-5 ft) thick (Hughes, 1980). The contact between this map unit and the underlying Madison Valley member is an angular unconformity in the northern Gallatin Valley.

Tscb Clarkston Basin member (Miocene)—Brownish gray and brownish yellow breccia or unconsolidated deposit with clast-supported or matrix-supported, fine to coarse, angular clasts of dominantly Greyson, Newland, and Spokane Formation, and Paleozoic quartzite and limestone. Paleozoic clasts are as large as boulder size. Coarse- and fine-grained beds may alternate. Cementation may be variable producing irregular surfaces.

Tscmp Madison Plateau member (Miocene: Clarendonian or Hemphillian)—Sheet deposit of moderately well-sorted and well-rounded clast-supported cobble conglomerate or gravel; color reflects maroon, gray, and brown calcium-carbonate-coated cobbles of dominantly Belt quartzite duraclasts. Archean metamorphic clasts are found only at the base as pebble size or smaller. The clasts were likely reworked from underlying Madison Valley member pebble conglomerates that contain Archean clasts. May correlate with clast-supported cobble conglomerate in Reese Creek area at base of Reese Creek member, from which an ash bed yielded a K/Ar date of 8.9 ± 0.4 Ma (Hughes, 1980; Lange and others, 1980). If so, the Madison Plateau member is likely youngest Clarendonian or oldest Hemphillian. West of the Madison River, unit appears to rest conformably on the Madison Valley member, but unconformably on the Dunbar Creek Member (Dunbar Creek Formation of Robinson, 1963) (Vuke and others, 2002). East of the Madison River, unit appears to rest conformably on the Madison Valley member and was involved in northeastward tilting and folding of underlying units with apparent conformity. Alternatively, unit may be younger than, and rest disconformably on the underlying Madison Valley member. Regardless, unit was deposited prior to northeastward tilting and folding of Tertiary strata. Thickness 6-10 m (20-30 ft).

Tsccc Cottonwood Canyon member (Miocene)—Brown conglomerate or gravel dominantly of rounded cobbles of Elkhorn Mountains Volcanics and Belt Supergroup quartzite, although almost exclusively derived from Elkhorn Mountains Volcanics west of the Boulder River. Conglomerate/gravel clasts include pebbles and small boulders in matrix of granules and coarse sand. Clasts are commonly stained with iron oxide. Conglomerate and gravel are in sheets and lenses interbedded with subordinate light gray, coarse-grained sandstone with conglomeratic stringers and floating granules, greenish brown or reddish brown, slightly bentonitic mudstone, tan siltstone, and micaceous, silty mudstone. Other than immediately west of the Negro Hollow uplift and east of Bull Mountain, unit contains abundant to occasional subrounded boulders of granitic rock with weathering rinds as much as ¼-inch thick. At one location south of the Negro Hollow uplift, extremely large boulders of granitic rock and Elkhorn Mountains Volcanics are present. Near Cottonwood Canyon this unit was called Ballard gravel by Aram (1979).

Near Cottonwood Canyon there appear to be two levels of deposits, both mapped as Cottonwood Canyon member. The upper deposit overlies a syncline in Cambrian units. It contains more clasts of intrusive rocks than the lower deposit, but similar to the lower deposit, does not contain Archean clasts. In the Cottonwood Canyon area, unit rests on LaHood Formation and Paleozoic rocks but does not contain clasts of either.

Barstovian fossils were found in the finer grained component of the unit between the Boulder River and the Negro Hollow uplift (Lofgren, 1985). Maximum exposed thickness as much as 100 m (330 ft).

Tsch Harrison member (Miocene and/or Pliocene?)—Grayish brown matrix-supported conglomerate or gravel with subrounded, texturally poorly sorted and disorganized clasts, ranging from pebbles to large boulders in iron oxide-stained matrix of granule and smaller size clasts. Clast composition dominantly Archean metamorphic rocks (mostly gneiss and schist) with subordinate quartzite, granitic rock including pegmatite, and Paleozoic limestone. Locally well-cemented, but generally unconsolidated. Linear distribution of unit extends along Antelope and Little Antelope Creeks to Sappington. Appears cut into older Tertiary units and in places rests on Archean and Paleozoic rocks. Linear geometry and presence of granitic boulders are similar to the Red Bluff Formation

near Norris, Montana, about ten miles to the southeast (Kellogg, 1994 and 1995). Maximum thickness as much as 60 m (200 ft).

Tscmv Madison Valley member (Miocene: Clarendonian and Barstovian)

Madison Bluffs, Madison Plateau, and Camp Creek Hills area. Pinkish tan or tan, tuffaceous silt or siltstone and marl interbedded with crossbedded, texturally immature, coarse-grained sandstone that contains lenses of pebble conglomerate or local cobble conglomerate ranging from matrix- to clast-supported, and from cemented to unconsolidated. Conglomerate clasts are dominantly Archean gneiss and extrusive volcanic rocks, with subordinate Belt rocks, and occasional Paleozoic limestone clasts. Sandstone has large root casts locally, and marl beds are typically full of small root casts. Several vitric ash beds are present throughout the unit. Opalized wood fragments are abundant in many conglomerate lenses. Conglomerate also contains relatively abundant disarticulated bones, bone fragments, and occasional teeth. Numerous articulated Barstovian fossils have been collected and studied from the fine-grained parts of this unit (Tabrum and Nichols, 2001; Douglass, 1899, 1901, 1903, 1907a, 1907b, 1908, 1909a, 1909b; Frick, 1937; Dorr, 1956; Sutton, 1977; Sutton and Korth, 1995; Evander, 1996), including the Anceney beds of Dorr (1956).

The contact between this unit and the underlying Dunbar Creek Member of the Renova Formation is sharp and locally appears unconformable as noted by Hackett and others (1960). In the northernmost part of the Madison Bluffs, the contact appears to be an angular unconformity. A relatively resistant bed at the base of the Madison Valley formation cuts down to nearly the level of the Climbing Arrow Member. In the southernmost part of the Madison Bluffs, calcic paleosols occur at the contact between the Madison Valley member and the underlying Dunbar Creek Member of the Renova Formation. Thickness 60 m (200 ft) throughout most of the Madison Bluffs area, but thickens to 90 m (300 ft) in the southern bluffs.

Dry Creek area. Grayish orange, crossbedded sandstone and pebble conglomerate interbedded with brownish orange tuffaceous siltstone and marl. Sandstone beds contain large lenses of dominantly pebble conglomerate with occasional cobble-size clasts or local cobble conglomerate. Conglomerates vary from matrix supported to clast supported. Unlike the clast composition in the Madison Bluffs and Camp Creek Hills areas, the clasts in the Menard area are dominantly Paleozoic limestone and Mesozoic sandstone, with subordinate andesitic volcanic rock and minor metamorphic rocks. Hughes (1980) interpreted the volcanic rock source as the Maudlow Basin where Livingston Group rocks are exposed. Near the Bridger Range, the clasts are dominantly Proterozoic LaHood Formation arkose, Paleozoic limestone, and quartz. Thickness near Menard about 30 m (100 ft). It is not known if the Madison Valley member thickens dramatically to the east, or if it has been down-faulted from the Bridger Range to Dry Creek. It is at a much higher elevation close to the Bridger Range than at Dry Creek.

Some of the coarse-grained beds of the Madison Valley member are shown on the map with a dotted pattern. In general, these are medium to coarse grained, fairly laterally persistent sandstone beds with numerous lenses of conglomerate and conglomeratic sandstone. Hughes (1981) describes these conglomeratic sandstones as “blanket-like” deposits in the Dry Creek area.

In local areas south of Manhattan and in the southeastern part of the map area, some of the conglomerate lenses contain dominantly cobble- rather than the more typical pebble-size clasts of other areas. These cobble conglomerates serve as caprocks with overlying fine-grained sediment stripped away. The cobble conglomerates have been interpreted as alluvial terrace deposits (Hackett and others, 1960) and pediment gravels (Mifflin, 1963), but were traced into the Madison Valley member in several places. They are shown as coarse-grained beds of the Madison Valley member. Gravel pits are numerous in the coarse-grained beds of the Madison Valley member.

Tscpb Parrot Bench member (Middle Miocene: Barstovian)—Brown conglomerate and gravel dominantly of subrounded to rounded cobbles of Elkhorn Mountains Volcanics and Belt Supergroup quartzite. Conglomerate/gravel clasts range from pebbles to small boulders in matrix of granules and coarse sand. Clasts are commonly stained with iron oxide. Resembles Cottonwood Canyon member which is also Barstovian, but has been included as part of the Parrot Bench map unit to the west (Vuke and others, 2004). Middle to Late Miocene Barstovian and Hemphillian fossils have been found in the Parrot Bench member immediately to the west of the map area on Parrot Bench (Kuenzi, 1966), but only the lower, probably Barstovian, part of the member is exposed in the map area.

Renova Formation (older than Hemingfordian unconformity)

Trnh Negro Hollow member (Early Miocene: Arikareean)—(Description modified from Lofgren, 1985). Dominantly poorly sorted, unconsolidated to moderately well cemented matrix-supported pebble conglomerate with granitic clasts dominant, and subordinate clasts of Elkhorn Mountains Volcanics, in a tuffaceous silt matrix. Conglomerate locally crossbedded. Lenses of clast-supported conglomerate generally at top of beds. Matrix-supported conglomerate grades laterally into, or is interbedded with, tabular beds of thinly laminated vitric silt and ash. Late Arikareean age determined from fossils (Lofgren, 1985). East of North Boulder River tuffaceous silt is dominant, and conglomerate is subordinate. Thickness about 275 m (900 ft).

Trdc Dunbar Creek Member (Eocene and Oligocene: Arikareean, Whitneyan, Orellan, and Chadronian)—The type section for the Dunbar Creek is within quadrangle in E½ sec. 7, T. 1 S., R. 2 E. (Robinson, 1963). Name derived from Dunbar Creek, a now obsolete name for a tributary of Mud Creek in the Three Forks area north of I-90.

Madison Bluffs and Camp Creek Hills. White to very light gray and light tan, tuffaceous planar-bedded siltstone and fine-grained sandstone with numerous tiny root casts, glassy ash beds, light gray tuffaceous limestone beds, and an interval of white diatomite beds near the top that extends the length of the Madison Bluffs. Distinguished from Madison Valley member by color, diatomite marker beds, and lack of significant conglomerate lenses. An $^{40}\text{Ar}/^{39}\text{Ar}$ date from the Dunbar Creek Formation in the Madison Bluffs on a single crystal feldspar yielded a range of 15-23 Ma with a weighted mean ~21.5 Ma, making it likely Arikareean North American Land Mammal age (upper Oligocene or lower Miocene). Hackett and others (1960) correlated the Dunbar Creek Member (their unit 1) in the Madison Bluffs and Camp Creek Hills with deposits near Menard in the Dry Creek area. Ostracode, gastropod, and fish fossils have been reported from the Dunbar Creek Formation in the Madison Bluffs (Blake, 1959; Dorr, 1956; Schneider, 1970). Thickness of unit about 45m (150 ft) in central Madison Bluffs, but may be thicker in the southern part of the Madison Bluffs where it is obscured by slope alluvium and colluvium. In the northernmost bluffs, it thins abruptly to about (15 m) 50 ft.

Calcic paleosols as described by Hanneman and others (1994) occur near road level in the area of Madison Buffalo Jump State Park (D. Hanneman, oral communication, 2002). They also occur where this unit is exposed in the northernmost and southernmost parts of the bluffs. Maximum exposed thickness about 45 m (150 ft).

Dry Creek area. White to very light gray and light tan, tuffaceous siltstone, fine-grained sandstone and mudstone, tuffaceous marl and limestone, and ash. Ostracode and gastropod fossils have been reported from this unit in the Dry Creek area (Hughes, 1980; Hackett and others, 1960; Verrall, 1955), and Klemme (1949) reported mammal fragments in this unit in the Dry Creek area. An ash in the Dry Creek area yielded a K/Ar date of 30.6±1.2 Ma (Whitneyan, Oligocene) (Hughes, 1980; Lange and others, 1980).

Willow Creek Reservoir area. Interbedded white, light gray, and light tan, very fine grained locally vuggy limestone and sandy limestone, white to light brown, tuffaceous, micaceous, calcareous siltstone and sandstone, and massive white ash. Contains silicified stems 1 cm in diameter with visible internal structure, and tubular root casts. Exposed thickness about 60 m (200 ft). Called Norwegian Creek beds by Feichtinger (1970).

Three Forks area. Very light gray to grayish yellow, thick-bedded, tuffaceous silt and siltstone with subordinate sand, sandstone, conglomerate, and rare bentonitic clay and very light gray limestone. North of the Jefferson River, unit includes many tongues of poorly sorted, poorly rounded carbonate-rich gravel not present south of the Jefferson River. The tuffaceous rocks and sediment are dominantly composed of fine volcanic ash. An Arikarean fossil, *Diceratherium armatum* Marsh, was found in a sandstone bed four miles south of Three Forks (Wood, 1933) in the lower Dunbar Creek Member. Also, an Arikarean beaver was described from the bluffs on the west side of the Madison River “about nine or ten miles south of Three Forks” (Douglass, 1901) in the Dunbar Creek Member. North of the Jefferson River the Dunbar Creek Member contains Chadronian (Eocene) fossils (Tabrum and others, 2001). North of the Jefferson River the contact between the Climbing Arrow and Dunbar Creek Members is conformable, and south of the Jefferson River, the contact is disconformable. Thickness as much as 90 m (300 ft) in the southeastern part of the map area.

- Trca Climbing Arrow Member (Eocene: Uintan, Duchesnean, Chadronian)**—Pale olive, light olive brown, and reddish brown, bentonitic, sandy clay and claystone that displays “popcorn” weathering; yellowish gray, coarse-grained, argillaceous sand and sandstone; and white, tuffaceous siltstone and fine-grained sandstone composed almost entirely of volcanic glass. Throughout, coarse sand grains are typically subangular to subrounded, and composed of quartz, feldspar, and biotite. A K/Ar date of 50.4±1 (Eocene) was obtained from an ash in the NE¼ NE¼ NE¼ sec. 11, T. 2 N., R. 1 W. (Lange and others, 1980). Type section is within quadrangle in W½ sec. 12, T. 1 N., R. 1 E. (Robinson, 1963). Thickness not less than 230 m (750 ft), but may be considerably more than 300 m (1,000 ft) (Robinson, 1963).
- Trmc Milligan Creek Member (Eocene)**—(Robinson, 1963): Light gray, fine-grained, tuffaceous, argillaceous limestone, marlstone, and calcareous mudstone units that interfinger with sandstone and conglomerate with rounded to subrounded clasts that are dominantly quartz and volcanic rock. Type section is within quadrangle in E½ sec. 11, NW¼ sec. 12, and SW¼ sec. 1, T. 1 N., R. 1 W.

Trrh **Red Hill member (Eocene: early Chadronian)**—Dominantly moderate red, reddish orange, and reddish brown kaolinitic mudstone with thin beds and lenses of pale olive gray or moderate red, very coarse grained, immature sandstone with clasts that include limestone, shiny black chert, clear and rose quartz, and red mudstone rip-up clasts; granule, pebble, and conglomerate lenses of similar composition in red mudstone matrix; and limestone breccia with red mudstone matrix, and red immature granitic sandstone. Unit typically weathers to red soil. Several vertebrate fossils found in member on Bull Mountain are early Chadronian (A. Tabrum, personal communication *in* Rothfuss, 2007). Member present on southern Bull Mountain and southern Doherty Mountain, and in Milligan Canyon and possibly Timber Canyon. Includes Sphinx conglomerate of Robinson (1963). Deposited in alluvial fan environments proximal to basin-margin uplifts (Rothfuss, 2007). Thickness variable; locally more than 90 m (300 ft) thick (Robinson, 1963).

Igneous rock

- Tan** **Andesite (Eocene)**—Medium gray, dominantly andesite porphyry, but also includes porphyritic latite, and latite porphyry. Many small phenocrysts of feldspar and mafic minerals are only slightly larger than the groundmass (Robinson, 1963). Exposed thickness as much as 110 m (360 ft).
- Trvi** **Rhyolite, vitrophyre (Eocene)**—White, pink, and gray, flow-banded, sparsely porphyritic rhyolite of dominantly brown cloudy glass and subordinate sanidine with sparse phenocrysts of altered biotite. Extensively brecciated near margin. Exposed thickness 260 m (850 ft).
- Tav** **Absaroka Volcanics (Eocene)**—Light brown-weathering, dark gray to black, basic, slightly porphyritic andesite with plagioclase, augite, and hypersthene phenocrysts in individual flows interlayered with stratified flow breccias. Exposed thickness 260 m (850 ft).
- Tba** **Basalt (Eocene?)**—(Robinson, 1963): Black, very dark gray, and grayish brown, fine-grained, intergranular basalt and basalt breccia flows along Cherry Creek fault. Commonly vesicular or flow banded. Vesicles in many places encrusted by yellow, fine-grained zeolite. Exposed thickness as much as 200 m (650 ft). Olivine basalt near Three Forks mostly compact but vesicular in some places. Homogeneous flow bands and minor flow breccia.
- TKa** **Andesite (Eocene or Late Cretaceous)**—(Dixon and Wolfgram, 1998): Porphyritic basaltic andesite on Bull Mountain, blocky to massive, 40-50 percent plagioclase phenocrysts that average 5 mm to 1 cm (0.2-2 in.) within a deep red to gray matrix. Silicified and argillic along contact with Mission Canyon Formation.
- TKl** **Latite (Eocene or Late Cretaceous)**—*Three Forks area* (Robinson, 1963): Medium gray to medium dark gray latite that weathers to steely purplish tones similar to the Elkhorn Mountains Volcanics. Latite has sparse and small phenocrysts of feldspar and pyroxene in a dense, nonvesicular holocrystalline matrix.
- TKda** **Dacite (Eocene or Upper Cretaceous)**—(Robinson, 1963): Medium light gray, with local yellowish brown or yellowish green splotches from oxidation of iron-bearing minerals; uniformly fine-grained. Dominantly dacite, but includes some quartz latite. Exposed thickness as much as 75 m (250 ft).

- TKfi Felsic intrusive rocks, undivided (Eocene or Late Cretaceous)**—*Red Mountain area* (Kellogg, 1995): Gray, pinkish gray, and purple porphyritic rhyolite or dacite with an aphanitic, cherty matrix. Plagioclase phenocrysts as long as 8 mm, but typically 3 mm. In most places flow banded. A columnar-jointed sill as thick as 25 m (82 ft) intruded the contact or lower part of the Flathead Formation southeast of Red Mountain.
- TKjb Jasperoid breccia (Eocene or Upper Cretaceous)**—Yellow, tan, and reddish brown, banded jasperoid that is almost entirely brecciated, and consists of angular clasts as much as 0.5 m (1.5 ft) thick. Exposed thickness as much as 75 m (250 ft).
- Ki Intrusive rocks, undivided**
- Kem Elkhorn Mountains Volcanics (Late Cretaceous)**—Dominantly dark gray, grayish black, greenish black, light gray, and very dusky reddish purple andesite porphyry, tuff, conglomerate, breccia, and minor flows; and subordinate basalt. Andesite porphyry displays variable texture and resistance. Conglomerates contain clasts that are rounded and as much as 50 cm (20 inches) wide in a coarse-grained sand matrix. Rests unconformably on units as old as the Lodgepole Limestone. As much as 2,750 m (9,000 ft) thick (Alexander, 1955).
- Kan Andesite (Late Cretaceous)**—Rocks with variable generally andesitic composition.
- Kdia Diabase (Late Cretaceous)**—*Horseshoe Hills* (Sayers, 1962): Highly altered, porphyritic augite diabase.
- Kg Granite (Late Cretaceous)**—*Tobacco Root Batholith* (Smith, 1970): Equigranular and fine- to medium-grained. Generally finer grained than other batholith rock types. Potassium feldspar megacrysts generally rare and relatively small, but occasionally as large as 3.0 cm. Plagioclase generally in euhedral crystals; quartz and potassium feldspars almost invariably anhedral. Potassium feldspars are poikilitic and include all other mineral types. Biotite, magnetite, apatite, sphene, and zircon accessory minerals.
- Kgd Granodiorite (Late Cretaceous)**—*Tobacco Root Batholith* (Smith, 1970): Dominantly fine-grained and equigranular to subordinate coarse-grained porphyries. Hornblende and biotite are visible megascopically. Potassium feldspar phenocrysts are typically smaller and less numerous than hornblende and biotite phenocrysts.
- Khd Hornblende diorite (Late Cretaceous)**—*Tobacco Root Batholith* (Smith, 1970): Hypidiomorphic granular to panidiomorphic fine- to medium-grained; 12-28% hornblende, 3-8%, quartz with accessory minerals biotite, pyroxene, sphene, magnetite, zircon, apatite, allanite, and rutile. (Kellogg, 1994): Dark gray, fine- to coarse-grained, equigranular to inequigranular, well indurated hornblende diorite and hornblende monzodiorite.
- Khto Hornblende tonolite (Late Cretaceous)**—*Tobacco Root Batholith* (Smith, 1970): Medium-grained and equigranular to porphyritic tonolite with 5-13% hornblende. Elongated ferromagnesian minerals emphasize flow structures. Potassium feldspars are poikilitic and quartz occurs as intergrowths with these and plagioclase. Granitoid texture is common and subhedral-euhedral crystals are completely interlocking. Quartz is interstitial and feldspars are zoned.

- Kto Tonalite (Late Cretaceous)**—*Tobacco Root Batholith* (Smith, 1970): Medium-grained and equigranular to porphyritic with pink potassium feldspar megacrysts in about 1/3 of the exposures. Platy elongated ferromagnesian minerals emphasize flow structures. Potassium feldspars are poikilitic and quartz occurs as intergrowths with ferromagnesian minerals and plagioclase. Granitoid texture is common and subhedral-euhedral crystals are completely interlocking. Quartz is interstitial and feldspars are zoned.
- Kqm Quartz monzonite or monzogranite (Late Cretaceous)**—*10N Pluton* (Robinson, 1963): Pink, pale grayish brown or moderate brown weathering, dominantly quartz monzonite and monzogranite, but also includes monzonite porphyry, diorite porphyry, and quartz latite porphyry. Strongly lineated. Exposed thickness 75 m (250 ft). *Tobacco Root Batholith* (Smith, 1970): Medium-grained (1.2-2.0 mm) ranging from equigranular to coarse porphyritic with 5 cm (2 in.) long potassium feldspar megacrysts. Most quartz monzonite at least slightly porphyritic.
- Ksk Skarn (Late Cretaceous)**—Contact metamorphosed Paleozoic carbonate rocks adjacent to the 10N Pluton north of Three Forks. Includes isolated, small igneous intrusive bodies (Michael Garverich, oral communication, 2012).
- Kmod Monzonite and diorite (Late Cretaceous)**—(Robinson, 1963): Porphyritic monzonitic and dioritic rocks of the 10N Pluton north of Three Forks
- Kse Sedan Formation (Upper Cretaceous)**—(Skipp and others, 1999):
Mudstone member: Greenish gray and brownish gray volcanoclastic mudstone, siltstone, sandstone, and minor interbedded conglomerate and altered vitric tuff.
Middle sandstone member: Olive green and dark greenish gray, volcanoclastic sandstone, conglomerate, mudflow conglomerate, and minor siltstone and mudstone.
Ash-flow tuff member: Pale yellowish green, light greenish gray, grayish red, and pale yellowish brown, welded to non-welded tuff and ash-flow tuff conglomerate, interbedded with volcanoclastic conglomerate, sandstone, mudstone, and porcellanite.
Lower sandstone member: Dark olive gray, greenish gray, and yellowish gray, volcanoclastic sandstone, siltstone, mudstone, altered crystal vitric tuff, minor hornblende dacite, and minor lignitic coal.
 Equivalent to Miner Creek and Cokedale Formations of the Livingston Group (Berg and others, 2000).
 Thickness of formation in map area about 915 m (3,000 ft).
- Ket Eagle and Telegraph Creek Formations, undivided**
Eagle Sandstone—Grayish orange to light olive gray, arkosic, cross-bedded fine- to medium-grained thin-bedded sandstone interbedded with siltstone. Contains carbonaceous siltstone, carbonaceous shale, and thin coal beds at base and top. Thickness about 75 m (250 ft).
Telegraph Creek Formation—Light olive gray to pale yellowish brown, thin-bedded to massive, very fine grained calcareous, arkosic sandstone and siltstone, interbedded with silty mudstone. Thickness about 75 m (250 ft).
- Kcot Cody through Thermopolis formations, undivided**
- Kcof Cody Shale and Frontier Formation, undivided (Upper Cretaceous)**

Cody Shale—Dark gray to brown mudstone interbedded with siltstone and very fine grained sandstone. Greenish gray, thin-bedded, glauconitic, fine-grained sandstone in middle of formation. Thickness about 150 m (500 ft).

Frontier Formation—Light brownish gray, fine- to coarse-grained thick-bedded to massive sandstone with subordinate siltstone. Dark gray to black, thin, chert-pebble conglomerate in some localities generally at bases of thick sandstone beds (Dyman and others, 1996; McMannis, 1952). Thickness about 170 m (550 ft).

Kbl Blackleaf Formation (Upper and Lower Cretaceous)—*Jefferson Canyon area* (McLane, 1971).

Upper: Gray, soft, ripple-marked, lithic sandstone and dark gray shale with marine fossils.

Basal: Brownish gray, very fine grained, crossbedded, resistant, quartz sandstone with quartz overgrowths. Exposed thickness of formation in Jefferson Canyon less than 30 m (100 ft).

Kmdt Muddy Sandstone and Thermopolis Shale, undivided

Muddy Sandstone: Interbedded micaceous, planar bedded or crossbedded sandstone and subordinate dark gray to black, fissile shale or mudstone.

Thermopolis Shale

Upper: Medium gray, fissile, micaceous, clayey shale with a few thin interbeds of siltstone (Dyman and others, 1996).

Basal: Yellowish gray- to pale olive-weathering, light gray, fine- to medium-grained sandstone with quartz overgrowths, crossbedded or ripple marked, clean, well-sorted quartz sandstone that may have interspersed limonite specks. Unconformably overlain by Elkhorn Mountains Volcanics in western part of map area. Thickness about 90-105 m (300-350 ft).

Kt Thermopolis Formation (Lower Cretaceous)

Upper: Dark gray to black, fissile shale to mudstone, that contains thin interbeds of micaceous, planar- or cross-bedded, lithic sandstone.

Middle: Medium gray, fissile, micaceous, clayey shale with a few thin interbeds of siltstone (Dyman and others, 1996).

Basal: Yellowish gray- to pale olive-weathering, light gray, fine- to medium-grained sandstone with quartz overgrowths, crossbedded or ripple marked, clean, well-sorted quartz sandstone that may have interspersed limonite specks. Unconformably overlain by Elkhorn Mountains Volcanics in western part of map. Thickness about 90-105 m (300-350 ft).

Kk Kootenai Formation

Upper: Light gray gastropod coquina or gastropod-rich limestone that may also contain charophyte and ostracode fossils. The gastropod limestone is not present in the Bridger Range (McMannis, 1952).

Middle: Variegated shale and mudstone, dominated by red, orange, and purple with subordinate light and medium gray colors, interbedded with light gray “salt-and-pepper” limonitic or non-limonitic, fine- to coarse-grained, poorly to well-sorted, massive or crossbedded, chert-rich, locally conglomeratic sandstone.

Basal: Light brown to yellowish gray “salt-and-pepper” conglomeratic, cross-bedded, chert-rich sandstone or conglomerate. Thickness about 120 m (400 ft).

Jme Morrison Formation and Ellis Group, undivided

- Jm Morrison Formation (Jurassic)**—Green, red, and gray variegated mudstone, shale, and siltstone with thin, interbedded yellowish brown to grayish orange, very fine grained sandstone and siltstone beds, and thin, gray limestone beds. Carbonaceous shale or coal at the top in some areas of the northern part of the quadrangle. Thickness about 105 m (350 ft).
- Je Ellis Group (Jurassic)**
- Jsw Swift Formation**—Grayish orange, calcareous, limonitic or glauconitic, crossbedded, coarse-grained, fossiliferous, quartz sandstone. Thickness about 20 m (65 ft).
- Rierdon Limestone**—Light gray, oölitic, fossiliferous limestone, and calcareous shale. Quartz and chert sand grains are interspersed throughout some of the limestone beds. Locally, quartz and chert clasts are granule or small pebble size. Thickness about 15 m (50 ft).
- Sawtooth Formation**
Upper: Yellowish brown, fossiliferous mudstone; thin-bedded fossiliferous carbonaceous siltstone and dolomite; and light gray, thin-bedded, fossiliferous limestone.
Lower: Gray to dark brown conglomeratic quartz and chert sandstone of variable thickness, with subangular to subrounded chert and light gray limestone pebbles. Formation not present in western part of map. Thickness 0-25 m (0-80 ft).
- Ʀd Dinwoody Formation (Triassic)**—Reddish brown to dark brown, fossiliferous, silty limestone with abundant *Lingula* brachiopods. Thickness 30-40 m (100-130 ft) (Elliott, 1998a).
- PPMps Phosphoria, Quadrant, and Snowcrest Range Group, undivided (Permian, Pennsylvanian, and Mississippian)**
- PPpq Phosphoria and Quadrant Formations, undivided (Permian and Pennsylvanian)**
- Pp Phosphoria Formation (Permian)**—Brown, to greenish brown, laminated or thin- to thick-bedded chert, yellow to yellowish orange sandstone and siltstone, greenish gray, medium- to coarse-grained, oölitic, phosphatic sandstone, and yellowish gray, dolomitic limestone. May also include conglomerate with well-rounded chert pebbles or cobbles. Thickness ranges from 30-60 m (100-200 ft).
- Pq Quadrant Formation (Pennsylvanian)**—Light gray, pinkish gray, and yellowish gray, medium- to thick-bedded, medium- to fine-grained, well-sorted, quartz sandstone with rounded clasts; cemented by quartz overgrowths. Very light gray, medium to thick dolomite beds may be present in the lowermost and uppermost parts, interbedded with quartz sandstone. Thickness variable, ranging from 15 m (50 ft) (McMannis, 1952) to 150 m (500 ft) (Robinson, 1963).
- PMsr Snowcrest Range Group, undivided (Pennsylvanian and Mississippian)**
- Conover Ranch Formation (Pennsylvanian and Mississippian)**—Red, blackish red, and pale red, irregularly bedded calcareous mudstone, siltstone, and sandstone, and gray dolomitic limestone. Not present locally in Three Forks area. Thickness 0-25 m (0-80 ft).
- Lombard Limestone (Mississippian)**—Dark gray, reddish gray, or black, fossiliferous shale, shaly limestone, and cherty limestone. Yellowish brown, silty, laminated to flaggy, finely crystalline limestone with interbeds of yellowish brown, highly calcareous siltstone, and thin interlaminae of dark gray chert. Thickness 0-130 m (0-425 ft).
- Kibbey Formation (Mississippian)**—Thin-bedded to massive, calcareous red, pink, and pale yellowish orange orthoquartzite, yellowish brown, silty, laminated to flaggy, finely

crystalline limestone with interbeds of yellowish brown, highly calcareous siltstone, and thin interlaminae of dark gray chert. Medium gray to yellowish orange fissile to hackly shale. Lower part has numerous thin interbeds of medium to dark gray, impure, calcareous, hard, quartzitic sandstone with grayish orange shale beneath each bed. Thickness 0-38 m (125 ft).

Mm Madison Group, undivided (Mississippian)

Mmc Mission Canyon Limestone (Mississippian)—Gray, microcrystalline, thick-bedded, locally fossiliferous limestone with abundant gray, black, olive black, and pale yellowish brown lentil-shaped or elongate chert nodules. Solution breccia and paleo-karst features are apparent in some areas. Variable thickness ranging from 115 m (375 ft) in the Gallatin Range (Tysdal, 1966) to 130 m (430 ft) in the Bridger Range (McMannis, 1952) to 460 m (1,510 ft) in the Three Forks area (Robinson, 1963).

MI Lodgepole Limestone (Mississippian)—Dark gray, thin-bedded, microcrystalline limestone, with yellowish brown and grayish orange thin partings and interbeds of calcareous mudstone. Basal Cottonwood Canyon Member black shale present in northwestern part of map area. Thickness ranges from 60-260 m (200-855 ft).

MDt Three Forks Formation (Mississippian and Devonian)—Sappington Member: Yellowish orange and yellowish gray thin- to thick-bedded, flaggy siltstone and fine-grained sandstone. May contain U-shaped trace fossils. Type section is in map area near Logan (Holland, 1952), and type locality is in the map area in Milligan Canyon near Sappington (Berry, 1943). Thickness 15-30 m (50-100 ft) (Sandberg, 1965).
Trident Member: Greenish gray, light olive gray, and yellowish gray calcareous to slightly calcareous fossiliferous clay shale with yellowish gray, dark yellowish orange, and medium gray dolomitic limestone, silty dolomite, and calcitic dolomite at the base. Massive bed of fossiliferous argillaceous limestone at the top. Type section is in map area about five miles northwest of Logan (Sandberg, 1965).
Logan Gulch Member: Yellowish gray and grayish red, argillaceous limestone or shale breccia that may be partly interbedded with dolomitic shale, dolomitic siltstone, and silty dolomite; yellowish gray, thin-bedded, contorted limestone; and red mudstone. Type section (Sandberg, 1962) is in the map area within the Three Forks type section (Sloss and Laird, 1947) north of the Gallatin River at Logan. Thickness ranges from 25-45 m (80-150 ft).

DЄjms Jefferson, Maywood, and Snowy Range Formations, undivided

Dj Jefferson Formation (Devonian)—Birdbear Member: Light brownish gray to medium gray, very finely crystalline to microcrystalline, sucrosic, partly pseudo-brecciated dolomite. Type section of member (Sandberg, 1965) is in map area within the type section of the Jefferson Formation (Sloss and Laird, 1947) on the north side of the Gallatin River near Logan. Thickness about 25 m (80 ft).
Lower member: Dark yellowish brown, brownish gray, medium dark gray, and light olive gray, finely crystalline, fetid dolomite, and calcitic dolomite. Thickness ranges from 145 to 200 m (475-655 ft).

DЄmr Maywood and Red Lion Formations, undivided

Maywood Formation (Devonian)—Grayish red, thin- to medium-bedded, aphanitic to very finely crystalline, dense to friable and shaly dolomite. Locally contains a bright pale

yellow sucrosic limestone bed in the upper part. Grayish red to yellowish orange calcareous siltstone at base in some places. Not present in the northern Gallatin Range (Tysdal, 1966). Thickness as much as 27 m (90 ft).

Red Lion Formation (Cambrian) (Western part of map area only, lateral equivalent of Snowy Range Formation)—Light pinkish gray, greenish gray, and light gray, sucrosic dolomite; grayish orange dolomite with silicified wavy shale partings; local flat pebble dolomite conglomerate. Thickness as much as 18 m (60 ft) in western part of map area.

D€msr Maywood and Snowy Range Formations, undivided

€sr Snowy Range Formation (Cambrian) (Eastern part of map area only, lateral equivalent of Red Lion Formation)—Sage Pebble Conglomerate Member: Medium light gray to light olive gray, thin- to medium-bedded, flat-pebble limestone conglomerate with subangular to subrounded clasts; and very finely crystalline to aphanitic, dense limestone with minor interbeds and interlaminae of greenish to red fissile shale and subordinate light grayish red, irregular-bedded or laminated siltstone, banded or mottled with yellowish orange. In the Bridger Range there is a persistent 1.22-7.62 m (4-25 ft) biostromal columnar limestone composed of calcareous fossil algae at the base of the member (McMannis, 1952). In the Horseshoe Hills, the biostromal columnar limestone is at the base up to the middle of the member, but is missing east of Nixon Gulch and northeast of Trident (Verrall, 1955). Thickness as much as 62 m (204 ft) in the Bridger Range (McMannis, 1952). Flat-pebble conglomerate is not present in the Gallatin Range, but a limestone member is present in the stratigraphic position of the Sage Pebble Conglomerate Member that also has a biostromal columnar limestone at its base (Tysdal, 1966).
Dry Creek Member: Light olive gray, grayish green, or bluish gray, fissile shale with interbedded pale orange to yellowish brown, calcareous, fine-grained sandstone and siltstone beds that commonly have scour bases. Thickness is irregular, ranging from 2-23 m (6-76 ft) (Tysdal, 1966; McMannis, 1952) in eastern part of map area.

€pi Pilgrim Limestone (Cambrian)—Light gray or bluish gray limestone or dolomite, typically with yellowish orange mottles. May be sandy or sucrosic; may contain intraformational flat-pebble conglomerate, or lenses of dark gray limestone or dolomite that are glauconitic, oölitic, and/or fossiliferous; weathers hackly. Thickness ranges from 60 m (200 ft) (eastern part of map area) to as much as 137 m (450 ft) (western part of map area).

€pm Park and Meagher Formations, undivided

€p Park Shale (Cambrian)—Grayish green and pale purple, fissile shale and silty shale. May contain a thin limestone bed or limestone flat-pebble conglomerate at the top, and thin interbeds of grayish red purple, coarsely crystalline, ferruginous limestone. Thickness ranges from 12-106 m (40-350 ft).

€m Meagher Limestone (Cambrian)—Light gray or bluish gray limestone, dolomite, or dolomitic limestone, with yellowish orange or moderate orange pink mottles; weathers hackly. Dominantly thick-bedded, but thin-bedded in part with siltstone partings. In the eastern part of the map area, may contain interbeds of greenish gray, micaceous, fissile shale. May contain oölitic beds with some oncolites and intraformational conglomerate. Thickness ranges from 90-167 m (300-550 ft).

- €w** **Wolsey Shale (Cambrian)**—Dominantly grayish green, but also grayish purple and grayish red purple, micaceous, fissile, wavy-bedded shale with trace fossils on many bedding surfaces. In western part of map area includes a greenish brown, carbonaceous, silty limestone in middle of unit. May be interbedded with thin quartzite beds at base. May contain trilobite casts and molds. In some areas plutonic rocks take the place of the Wolsey Shale (Robinson, 1963). Thickness ranges from 12 m (40 ft) in the Gallatin Range (Tysdal, 1966) to as much as 120 m (400 ft) (western part of map).
- €f** **Flathead Formation (Cambrian)**—Very light gray, pinkish gray, or light brownish gray quartzose sandstone or orthoquartzite and well-cemented granule to pebble conglomerate. May be massive or crossbedded and contain subordinate grayish green, grayish purple or grayish red purple, micaceous, fissile shale beds. In some areas in the central part of the map area, the Flathead Formation is missing. Basal unconformity places the Flathead on Archean rocks south of the Willow Creek fault, and on Proterozoic rocks north of the fault. Where present, thickness ranges from 12-45 m (40-150 ft).
- Yg** **Greyson Formation (Mesoproterozoic)**—Greenish gray and yellowish brown siltite and fine-grained quartzite with subordinate limestone (Robinson, 1967). Exposed thickness 200 m (650 ft).
- Yn** **Newland Formation (Mesoproterozoic)**—Gray and yellowish brown, calcareous siltite, limestone, and fine-grained quartzite (Robinson, 1967). Exposed thickness 150 m (490 ft).
- Yla** **LaHood Formation, undivided (Mesoproterozoic)**—Dark gray, dark brownish green, and locally reddish brown, arkosic boulder to granule conglomerate, arkose, arkosic siltite, arkosic argillite, and some impure carbonate beds. Clasts composed primarily of various Archean metamorphic, and igneous rocks; matrix light olive gray mudstone. Type section in map area near LaHood Park (Alexander, 1955; McMannis, 1963). Clast size as much as 3.6 m (12 ft) wide in local areas, but size decreases dramatically in lobate patterns that fan northward. Clasts are angular to subround.
Shale facies (Greyson Shale of Alexander, 1955): Black, brown, or purple, fissile, to non-fissile, carbonaceous or silty shale, interbedded with gray siltite and argillite, coarse- to fine-grained arkose, and thin carbonate beds.
Carbonate facies (Newland Limestone of Alexander, 1955): Dark gray to light gray, thin-bedded to laminated limestone, and olive-gray dolomite; locally contains algal structures (Verrall, 1955).

LaHood Formation facies descriptions that follow were taken from an unpublished report by Tor Nilsen for the Golden Sunlight Mines, Inc.

Ylaf **LaHood Formation, alluvial-fan and fan-delta facies**

The best observed exposures are in the Huller Spring area and vicinity of Lewis and Clark Caverns.

Debris-flow component: Reddish to purplish weathering, massively bedded sandstone that lacks internal current-formed sedimentary structures. Displays reverse grading, matrix-supported angular clasts, poor sorting of both matrix and clasts, and isotropic fabrics.

Stratified component: Grayish-weathering, planar-bedded and low-angle crossbedded sandstone with local shale drapes and interbeds, moderately to well sorted, subangular to rounded clasts and rip-up clasts of shale; local channel-form geometries, and fining-

upward cycles at several localities. Shale interbeds characteristically weather to brown and olive brown and locally exhibit mudcracks, wave ripples, and parallel lamination. At Huller Spring, large angular blocks of migmatite, and angular to rounded clasts of quartzofeldspathic gneiss characterize the basal part of the facies. Overlying deposits consist chiefly of massive, pebbly, arkosic sandstone, interbedded with crossbedded and planar bedded pebbly arkosic sandstone, and minor amounts of argillite.

Facies fines to north, and has larger component of stratified beds with better rounded and sorted clasts than to the south. Unit locally represents fan-delta deposits into small lakes.

Ylsh LaHood Formation, shelf facies

Crops out south of Jefferson River along western border of map.
Massive to parallel bedded, pebbly sandstone; stratigraphically between submarine-canyon and alluvial fan facies.

Ylsl LaHood Formation, slope facies

Crops out in the Cave Fault area.
Characterized by abundant synsedimentary slumps. Consists dominantly of mudstone with thin siltstone turbidites and graded non-turbidite intra-flow deposits. Highly tectonized because of the lack of competent beds and appears to consist mostly of sheared and deformed argillite.

Ylsc LaHood Formation, submarine-canyon facies

Crops out in Lewis and Clark Caverns State Park area, where it appears to have cut directly into the alluvial fan/fan-delta facies.
Boulder to pebble conglomerate and conglomeratic sandstone that is poorly stratified and highly channelized. Canyons relatively small—less than 2 km (6,560 ft) wide—and are cut into shallow marine units, so exposures are principally of the upper parts of the canyons.

Ylis LaHood Formation, inner submarine-fan facies

Crops out in LaHood Canyon and in the Cave Fault area.
Channel axis component: Boulder to pebble conglomerate and pebbly sandstone with thin shale interbeds.
Channel-margin component: Argillite with thin interbeds of turbidite siltstone and very fine grained sandstone.

Ylms LaHood Formation, middle submarine-fan facies

Crops out east of the Golden Sunlight Mine, in LaHood Canyon, and in the Cave Fault area.
Characteristic repetitive fining- and thinning-upward sequences with channelized bases and interbedded, thin-bedded overbank turbidites. Overbank turbidites consist of thin, but coarse-grained, laterally discontinuous deposits with abundant small-scale slump folds that generally indicate directions of slumping away from channel axes. Locally contains large boulders of marble and limestone in LaHood Canyon and Cave fault areas. Apron-like fan geometries along the steep Proterozoic Willow Creek fault promoted transport of boulders into the middle-fan channels.

Ylos LaHood Formation, outer submarine-fan facies

Crops out east of the Golden Sunlight Mine, north and east of Cardwell, and in LaHood Canyon.

Typically well exposed because of the abundant resistant sandstone beds that make up stacked, thickening and coarsening-upward sequences. Characteristic Bouma sequences and abundant sole markings. Fan-fringe deposits commonly contain thinner and finer-grained beds organized into thinner coarsening-upward sequences with lower sandstone-to-shale ratios than central parts of fans. Individual beds are typically planar, well-graded, and separated by shale intervals. Slurried beds, in which the Bouma B division contains shale rip-up clasts that are locally very large, are common in many sections and can be mistaken for slump folds.

Ylbp LaHood Formation, basin-plain facies

Crops out south and east of the Golden Sunlight Mine.

Generally poorly exposed because consists mostly of argillite with thinly interbedded, laterally continuous and planar-bedded, distal turbidites. Abundant sole markings on relatively thick beds of sandstone. Generally has a lower sandstone-to-shale ratio, thinner beds of turbidite sandstone, and lacks thickening- and coarsening-upward sequences compared to the outer fan facies. Facies generally has lower sandstone-to-shale ratio and thinner and finer-grained beds of sandstone to the north.

The following two units are found in Precambrian basement rock in the map area. Although important, they could not be shown well at the map scale.

Diabase (Neoproterozoic)—Numerous dikes that cut across Archean rocks, but do not extend into younger rocks. (Schmidt and Garihan, 1986b).

Pegmatite and quartz veins (Late Cretaceous, and Paleoproterozoic)—White to pink, coarse-grained to very coarse grained, massive and foliated dikes and sills composed mostly of potassium feldspar, quartz, plagioclase, muscovite, and rarely, biotite. In some places grades into quartz veins that are white, massive quartz in lenticular, generally discordant veins and irregular pods. Very widespread in areas of Archean rocks.

Xsp Spuhler Metamorphic Series (Paleoproterozoic)—Light gray to reddish brown, purplish brown, or deep golden brown anthophyllite-gedrite gneiss and subordinate schist. Weathers to dark brown or blackish brown.

Xg Granite (Paleoproterozoic?)—Possibly Cretaceous.

XAah Amphibolite and hornblende gneiss (Paleoproterozoic and Archean?)—Gray to black, medium-grained, hypidiomorphic, equigranular, moderately foliated to well-foliated hornblende-plagioclase gneiss and amphibolite.

XAif Banded iron formation (Paleoproterozoic and Archean?)—Dark reddish brown to orange brown, massive to layered quartz-hematite rock locally containing abundant quartz veins; limonitic, especially along fractures.

XAq Quartzite (Paleoproterozoic and Archean?)—White, gray, and brown, medium- to coarse-grained, inequigranular, moderately foliated to massive quartzite.

XAqa Quartzite and amphibolite (Paleoproterozoic and Archean?)—Interlayered white quartzite and amphibolite.

- XAqfg Quartzofeldspathic gneiss (Paleoproterozoic and Archean?)**—(Vitaliano and Cordua, 1979): Includes plagioclase-microcline-quartz biotite (“granitic”) gneiss, plagioclase-quartz-biotite (“tonalitic”) gneiss, banded biotite gneiss, aluminous gneiss and schist, gedrite gneiss, and garnet gneiss.
- Plagioclase-microcline-quartz-biotite gneiss.* Light gray to light pinkish gray, medium-grained, weakly to moderately foliated gneiss ranging from granodiorite to syenogranite.
- Plagioclase-quartz-biotite gneiss.* Gray, medium grained, inequigranular, weakly to moderately foliated, tonalitic gneiss. Includes some trondhjemitic and granodioritic gneiss.
- Banded biotite gneiss.* White, light gray, dark gray, and black, medium-grained, well-foliated, inequigranular, tonalitic to quartz monzonitic gneiss, commonly migmatitic.
- Aluminous schist and gneiss.* Gray to dark brownish gray, medium-grained, inequigranular, generally well foliated, commonly micaceous gneiss and schist containing aluminosilicate minerals.
- Gedrite gneiss.* Brown to grayish brown, moderately well foliated, medium-grained, gedrite gneiss. Generally occurs in small lenses and concordant layers in other Archean rocks.
- Garnet gneiss.* Highly garnetiferous assemblage of various colors that includes biotite-garnet schist, sillimanite-garnet schist, garnetiferous quartzite, quartzite, garnetiferous quartzofeldspathic gneiss, corundum gneiss, gedrite schist, cummingtonite schist, and garnetiferous amphibolite.
- XAum Ultramafic rock (Paleoproterozoic and Archean?)**—Includes mafic to intermediate gneiss, hornblende-plagioclase gneiss, amphibolite, granulite, and intrusive metabasite.

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*Current name: Montana Tech of The University of Montana